

# West-European terrestrial molluscs assemblages of isotopic stage 11 (Middle Pleistocene): climatic implications

Denis-Didier Rousseau<sup>a</sup>, Jean-Jacques Puisségur<sup>b</sup> and François Lécolle<sup>c</sup>

<sup>a</sup>*Lamont-Doherty Geological Observatory of Columbia University, Palisades, N.Y. 10964, USA*

<sup>b</sup>*URA CNRS 157, Centre des Sciences de la Terre, 6 Bd Gabriel F-21100 Dijon, France*

<sup>c</sup>*Groupe Seine, Centre de Géomorphologie du CNRS, rue des Tilleuls, F-14000 Caen, France*

(Received May 16, 1991; revised and accepted October 10, 1991)

## ABSTRACT

Rousseau, D.-D., Puisségur, J.-J. and Lécolle, F., 1992. West-European terrestrial molluscs assemblages of isotopic stage 11 (Middle Pleistocene): Climatic implications. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 92: 15–29.

Tufa deposits in the Seine and the Avelasse Valleys have yielded numerous molluscs assemblages belonging to the *Lyrodiscus* Biome. These malacofaunas, primarily indicative of forest which developed under a warm climate, have no present-day counterpart. The biostratigraphical study of several localities based on new investigations at Vernon (Seine Valley) and Arrest (Avelasse Valley) and a reexamination of the faunas from St Pierre-lès-Elbeuf suggests the allocation of a Holsteinian age. This is supported by U/Th dates for Vernon and the stratigraphy of the loess sequence of St Pierre-lès-Elbeuf. Similar assemblages yielded by tufas of the same age in England and Germany indicate that the *Lyrodiscus* Biome was widespread. The warm climate recorded by the molluscs of the tufa deposits is in agreement with pollen and marine analyses for the isotopic stage 11 (362–423 kyr B.P.). In this way, the *Lyrodiscus* Biome appears as an interesting marker for the continental Quaternary in Western Europe. If the occurrence of these particular molluscan assemblages is due to climate, refuges from the declined Tertiary biota, from which i.e. *R. (Lyrodiscus)* could have emigrated, also needed to exist.

## Introduction

Among biological remains which occur in continental deposits, pollen are frequently used to reconstruct paleoenvironments. Numerous studies demonstrate that climatic changes occurred at roughly the same times in the oceans and on the continents (Peteet et al., 1990; Rossignol-Strick, 1983; Woillard, 1978; Van der Hammen et al., 1972). In other geological contexts, molluscs also provide good ecological data which complement the botanical data (Kerney, 1963; Lozek, 1964; Puisségur, 1976). In the Achenheim loess series (Rousseau and Puisségur, 1990a), variations in the components of the molluscan associations are similar to the isotopic  $\delta^{18}\text{O}$  curve provided by SPEC-

MAP (Imbrie et al., 1984). Here we report the results of a malacological investigation made in terrestrial tufas, from the Seine and the Somme valleys in northern France (Fig. 1). A comparison is made with continental pollen and marine records.

## New malacological records

### *Vernon*

The Vernon tufa is a complex deposit, located in the Seine Valley, about 70 km west of Paris (Fig. 1) which shows two different stages of development (Lécolle et al., 1990) (Fig. 2):

(a) A thick travertine, well indurated in places or showing alternating discontinuous smooth-indurated levels, constitutes the base of the upper part. The travertine is one to several meters thick,

Correspondence to: D.-D. Rousseau, Lamont-Doherty Geological Observatory of Columbia University, Palisades, N.Y. 10964, USA.

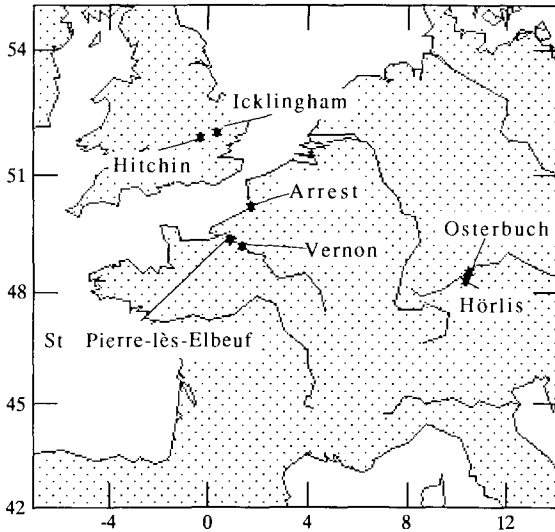


Fig. 1. Location of the different sites discussed in the text.

and consists of a so-called slab, rich in molded plant remains such as roots, branches or leaves. A two meter or so thick, smooth tufa overlays the travertine and is locally indurated. Sediment was taken for malacological investigation from this smooth unit. All these horizons are light in color. The morphology and position of the bed-rock are unknown.

(b) The lower part is dissociated from the previ-

ous one because it corresponds to the colluviation of upper part of the tufa and its rubefaction. According to a geological and geomorphological study (Lécolle, 1989), this second part corresponds to another time of deposition, clearly distinct from the upper one.

Determination of the species, present from a previous sampling in the upper non-cemented tufa, indicates that the assemblages are unique in their composition (Table 1). They do not correspond to any classical association as determined by Puissegur (1976) for northern France, but seem to be the western equivalent of the eastern European Banatica fauna (Rousseau, 1992). Four taxa are extinct: *Retinella (Lyrodiscus) skertchlyi* (Rousseau and Puissegur, 1990b), *Aegopinella bourdieri* (Rousseau and Puissegur, 1989), *Succinea joinvillensis* and *Catinella antiqua*. Six species no longer live in northern France: *Leiostylia anglica*, endemic in the British Islands, *Hygromia cinctella* present in Southern Europe, and *Discus ruderatus*, *Acicula polita*, *Aegopis verticillus* and *Ruthenica filograna* from Central Europe (Kerney et al., 1983) (Fig. 3). The other species still live in northern France, and a few aquatic taxa are also present.

In addition to this apparent heterogeneity of the assemblage, the occurrence of *Retinella (Lyrodiscus) skertchlyi* is another peculiarity. This species

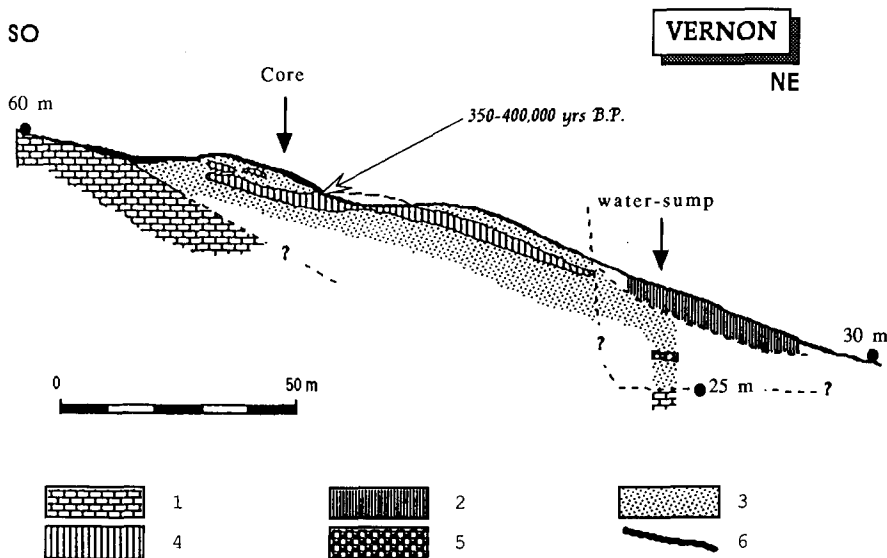


Fig. 2. Geomorphological and stratigraphical interpretation of the tuffaceous formation in Vernon. 1= substratum (chalk); 2= rubefaction; 3= smooth facies tufa; 4= travertine; 5= "sugar-plum" facies tufa; 6= colluvium (from Lécolle et al., 1990 modified).

TABLE 1

Vernon, first malacological investigations. List of different species recognized in the samples and individuals counted.

Species \ Levels	1	2	3	4	5	/ Ecol. Groups
<i>Acanthinula aculeata</i>	23	31	6	4	4	
<i>Acicula polita</i>	3			1		
<i>Azeca goodalli</i>	2	15	2	3	3	
<i>Aegopinella nitidula</i>	47	90	12	33	35	
<i>Aegopinella pura</i>	38	96	15	15	18	
<i>Aegopinella bourdieri</i>	1					
<i>R. (Lyrodiscus) skertchlyi</i>	1	23	4	18	30	
<i>Helicodonta obvoluta</i>	22	5	3	7	5	1
<i>Clausilia bidentata</i>	20	8	1	1	13	
<i>Cochlodina laminata</i>	7	6	3		3	
<i>Perforatella incarnata</i>		8	3	10	8	
<i>Orcula doliolum</i>		25	2	4	45	
<i>Aegopsis verticillus</i>		1			1	
<i>Ena montana</i>				3	2	
<i>Ruthenica filograna</i>				19	36	
<i>Discus rudерatus</i>					6	
<i>Discus rotundatus</i>	53	95	24	70	1	
<i>Acicula dupuyi</i>	11					
<i>Cepaea sp.</i>	35	50	18	15	50	
<i>Bradybaena fruticum</i>	4	15	6	4	7	
<i>Hygromia cinctella</i>	12					
<i>Pomatias elegans</i>	45	60	18			2
<i>Laciniaria biplicata</i>	2	4	1	10	30	
<i>Vitrea crystallina</i>	67	130	20			
<i>Arianta arbustorum</i>		12	4	5	6	
<i>Macrogastra ventricosa</i>	9	1		1	1	3
<i>Cecilioides acicula</i>	4					4
<i>Vallonia costata</i>			1	2	29	5
<i>Trichia hispida</i>	6	43	6	40	70	
<i>Helicigona lapicida</i>	2	4	4	1	2	
<i>Euconulus fulvus</i>	1	2				
<i>Punctum pygmaeum</i>	2					
<i>Vitrina pellucida</i>	6	6	1	4		7
<i>Oxychilus alliarius</i>	21					
<i>Clausilia parvula</i>	7	5		4		
<i>Vitrea contracta</i>	23	49	8	20	3	
<i>Cochlicopa lubrica</i>		31	7	14	7	
<i>Limax sp.</i>	10	16	4	2	25	7'
<i>Carychium tridentatum</i>	2	20	2		3	
<i>Succinea joinvillensis</i>	13	4	3		5	
<i>Succinea oblonga</i>	1					8
<i>Catinella antiqua</i>		1	1			
<i>Columella edentula</i>	1					
<i>Leiostylia anglica</i>		3	1	1		
<i>Succinea putris</i>		10	3	6		9
<i>Zonitoides nitidus</i>				1		
<i>Belgrandia sp.</i>	1	2		1		
<i>Galba truncatula</i>		7	1	2		10
<i>Pisidium sp.</i>		1				

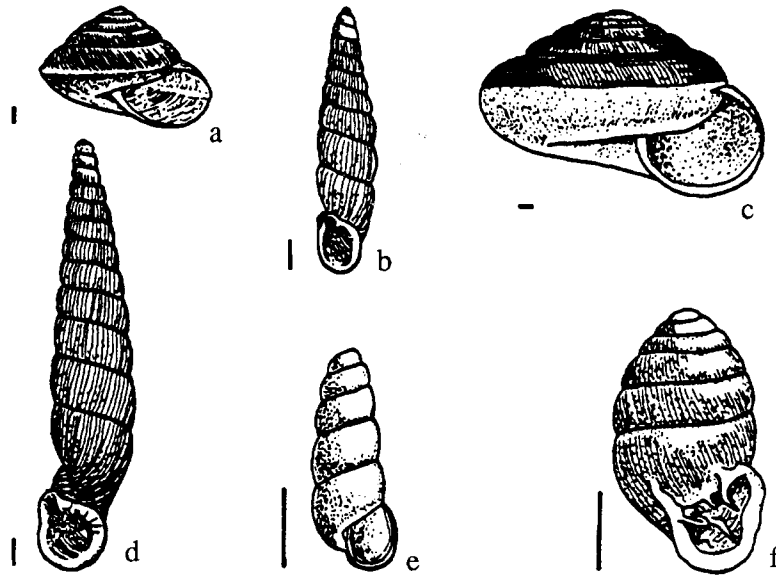


Fig. 3. Some of the allochthonous species, according to their present distribution, occurring in tufa assemblages. a. *Hygromia cinctella*; b. *Ruthenica filograna*; c. *Aegopis verticillus*; d. *Laminifera pauli*; e. *Acicula polita*; f. *Leiostyla anglica*. The scale represents 1 mm.

has related endemic species in the Canary archipelago (Kerney, 1976; Rousseau and Puisségur, 1990b).

The assemblages show a high dominance of forest species (Fig. 4a, b). This implies a forest environment that developed under a climate with warmer and drier summers than today in the same area. The occurrence of fossil leaves and seeds, including *Ficus*, which today has a more southern natural distribution, confirms the climatic interpretation.

A new investigation provides more detailed data

from these pulverent deposits (Fig. 2 and 5). Above the travertine, 9 samples were taken corresponding to different recognized horizons according to color and induration of the sediment. Two additional samples were taken to control the base of the core and correlated on the basis of lithology, with the first two samples of the main core.

Three kinds of deposits, from bottom to top (Fig. 5), overlay the travertine:

- a relatively non-cemented tufa, with weakly indurated levels of various colors (–3.6 to –2 m),
- deposits with a loamy facies which seem to

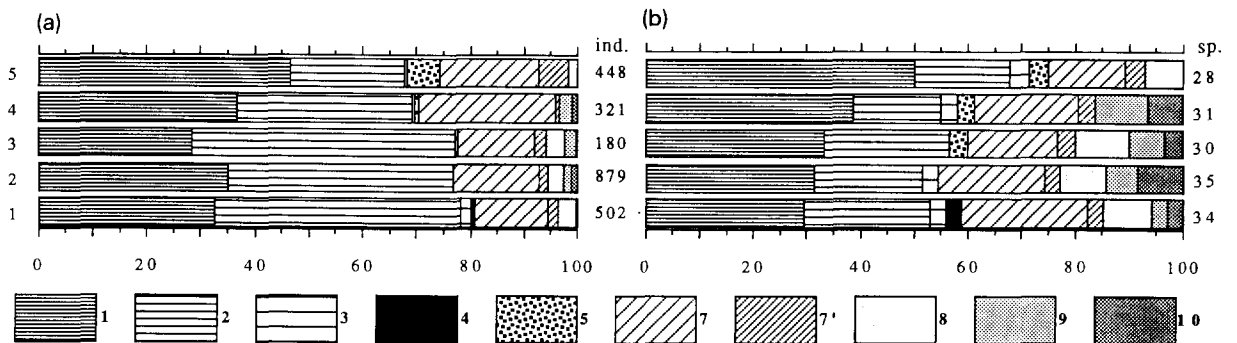


Fig. 4. Vernon. First malacological investigations. Variation of the percentage in (a) individuals (*ind.*) and (b) species (*sp.*) within each ecological group, according to Puisségur (1976). Signification of the groups: 1 = forest, 2 = open forest, 3 = damp forest, 4 = steppe, 5 = open ground, 7 = mesophilous taxa, 7' = slugs, 8 = hygrophilous taxa, 9 = paludal, 10 = fresh water.

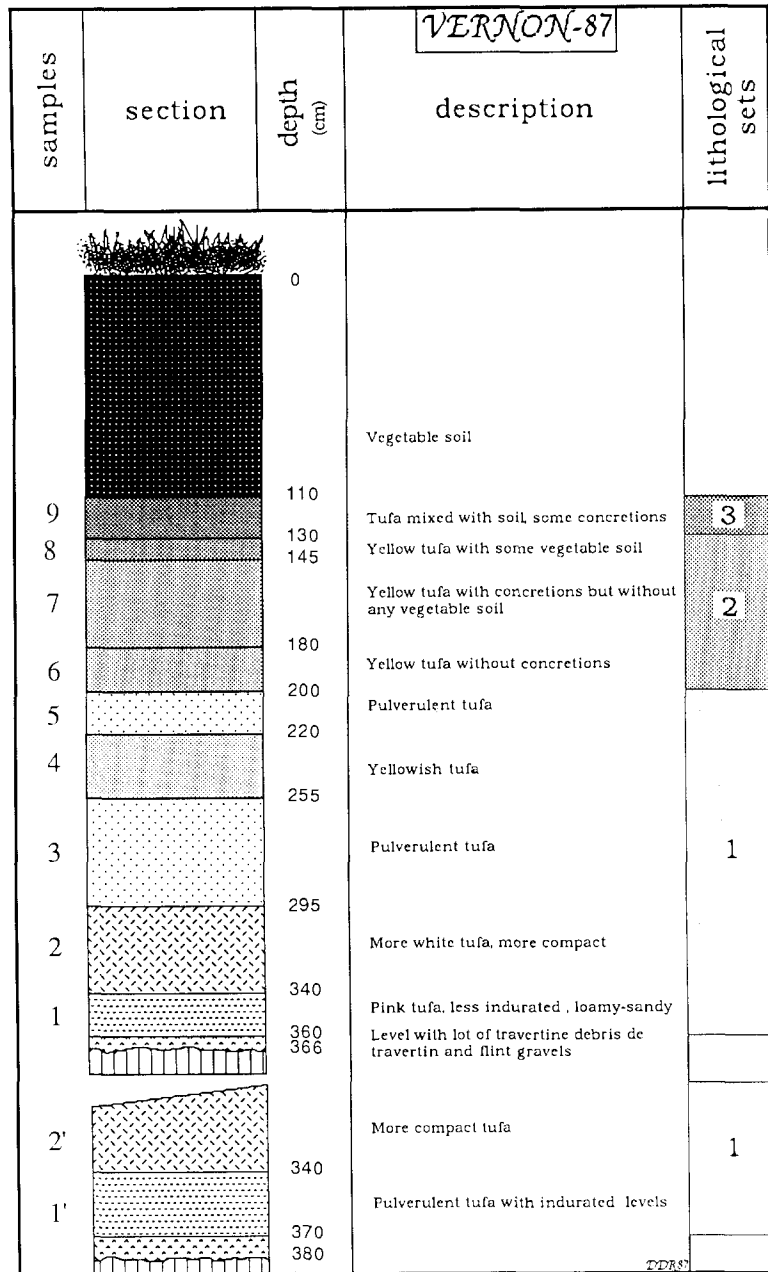


Fig. 5. Vernon. New malacological investigations. Location of the different malacological samples and schematic lithological description of the levels. 1-3 correspond to the different groups recognized.

correspond to a degradation of the tufa (-2 to -1.3 m)

—a composite of tufa and soil under the modern soil.

The list of the recognized species shows high

diversity and abundance (Table 2). Moreover, the variations in the total number of individuals occurring in each ecological group, sensu Puisségur (1976), is in agreement with the three previously defined units based on lithology (Fig. 6).

TABLE 2

Vernon, new malacological investigations. List of different species recognized in the samples and individuals counted.

Species \ Levels	9	8	7	6	5	4	3	2	1	2'	1'	/ Ecol. Groups
<i>Aegopinella nitidula</i>	3		4	4	34	25	35	12	20	18	6	
<i>Aegopinella pura</i>		1		1	14	21	12	3	2	6	3	
<i>Helicodonta obvoluta</i>		1	3	2	31	10	8	6	6	8	8	
<i>Ena obscura</i>			1									
<i>Azeka goodalli</i>			1	2	2	2	5		1	3		
<i>Aegopinella bourdieri</i>					5	4						
<i>Cochlodina laminata</i>					2	3	6	2		1	1	1
<i>Perforatella incarnata</i>					3	1		1		1		
<i>Acanthinula aculeata</i>					3	1	4	4	4	5	1	
<i>Aegopsis verticillus</i>					3	5	1		2		1	
<i>Clausilia bidentata</i>					4	5	3	2	4	2		
<i>Pomatias elegans</i>	20	4	3		9	30	17	18	9	5	4	
<i>Discus rotundatus</i>	7	1	3	5	17	22	8	31	12	8	1	
<i>Cepaea sp.</i>	7	3	2	5	35	35	30	28	27	10	16	
<i>Hygromia cinctella</i>			1	2	6	2	1	2	1	1	2	
<i>Laciniaria biplicata</i>			1		2	2						2
<i>Arianta-Bradybaena</i>					1	1	2		1		1	
<i>Acicula dupuyi</i>						3	1			15		
<i>Vitrea crystallina</i>							1		7	1	2	
<i>Acicula polita</i>										1		
<i>Macrogastra ventricosa</i>					1	2	2	1	1		1	3
<i>Cecilioides acicula</i>	37	6	2	3								
<i>Helicella sp.</i>	12	7	3	1								4
<i>Vallonia costata</i>	200	16	15	2								
<i>Vallonia pulchella</i>	50	3										5
<i>Pupilla muscorum</i>	3	1	2	1								
<i>Helix aspersa</i>		1	1	1								
<i>Trichia hispida</i>	213	16	17	1	3	3	5			1		
<i>Cochlicopa lubrica</i>	43	1	1									
<i>Oxychilus sp.</i>	5	1	3	1	3							
<i>Nesovitrea hammonis</i>	5											
<i>Vitrea contracta</i>	1	1	1	2	7	3	1	1	1	1		7
<i>Clausilia parvula</i>	3			1								
<i>Helicigona lapicida</i>					4	4	2	1	1	1	1	
<i>Vitrina pellucida</i>				1	13	3	3	2	2	2		
<i>Punctum pygmaeum</i>									1			
<i>Limax sp.</i>	13		1		1	2	2	1		1	3	7'
<i>Succinea oblonga</i>	71	3	1	2							1	
<i>Carychium tridentatum</i>									1			8
<i>Succinea putris</i>	12					2						
<i>Zonitoides nitidus</i>	10		2	2	6	3	3	4	4	5		9
<i>Carychium minimum</i>	3											
<i>Planorbis planorbis</i>	5											
<i>Galba truncatula</i>	4											
<i>Belgrandia sp.</i>						1	1					10
<i>Ancylus fluviatilis</i>							1					

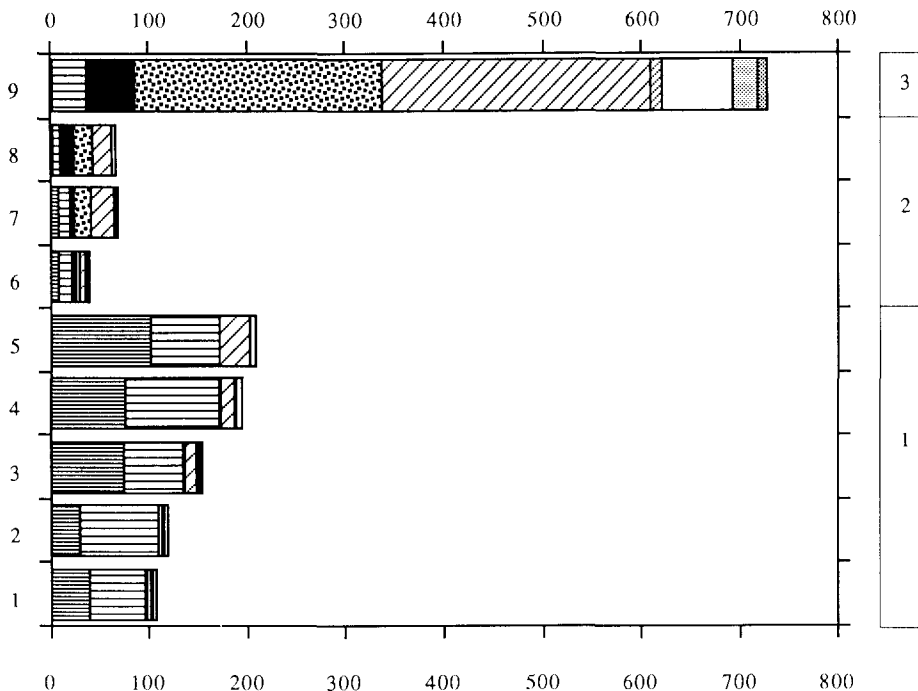


Fig. 6. Vernon. New malacological investigations. Variation of individual numbers in each ecological group (see legends in Fig. 4).

An increasing number of individuals from bottom to top and a high percentage of forest taxa characterize the malacofauna of the first unit. The second unit shows a generally decreasing number of individuals, a decrease in forest species, and the occurrence of steppe and open ground taxa. Finally the number of individuals increases sharply again in the top sample. This increase does not correspond to a bias in the sampling as this is a 20 cm sample taken under conditions roughly similar to the others.

This tripartition occurs when the different assemblages are compared, using species and individuals spectra of the following ecological groups: (1) forest, (2) semi-forest, (3) moist forest, (4) steppe, (5) open ground, (7) mesophilous, (7)' slugs, (8) hygrophilous, (9) paludal and (10) aquatic species (Fig. 7a, b). These ecological groups are defined according to the ecology of each mollusc species (Lozek, 1964; Puisségur, 1976).

The first unit reflects the high dominance of forest taxa, especially in the number of individuals, and has a small aquatic component (three individuals in samples 3 and 4 (fig. 7). However, moisture is indicated by the occurrence of *Macrogastra*

*ventricosa* (Table 2). The lack of aquatic species agrees with the occurrence of a non-permanent spring, an interpretation previously suggested by Puisségur (Puisségur in Lécalle, 1989). At the top of this unit (sample 5; Table 2), the proportion of mesophilous species increases. The species *Helicigona lapicida*, which has a preference for a slight forest covering, does not occur above sample 5. This development agrees with the general evolution of the malacofaunas.

Variations in group 7 permit the definition of the second unit (Fig. 7), which represents a more open environment. In this unit the progressive decrease of the forest taxa occurs from bottom to top (fig. 7), corresponding to increasing steppe and open ground species. The forest component gradually decreases after sample 5 and is replaced by a grassland in which still occur some trees or bushes.

The last sample, corresponding to the third unit, shows a completely open environment in which few bushes remain.

Characteristic taxa, *Hygromia cinctella*, *A. bourdieri*, *A. verticillus* and *A. polita* (Fig. 3), occur only in the first unit (Table 2) confirming its interglacial allocation, followed by unit two repre-

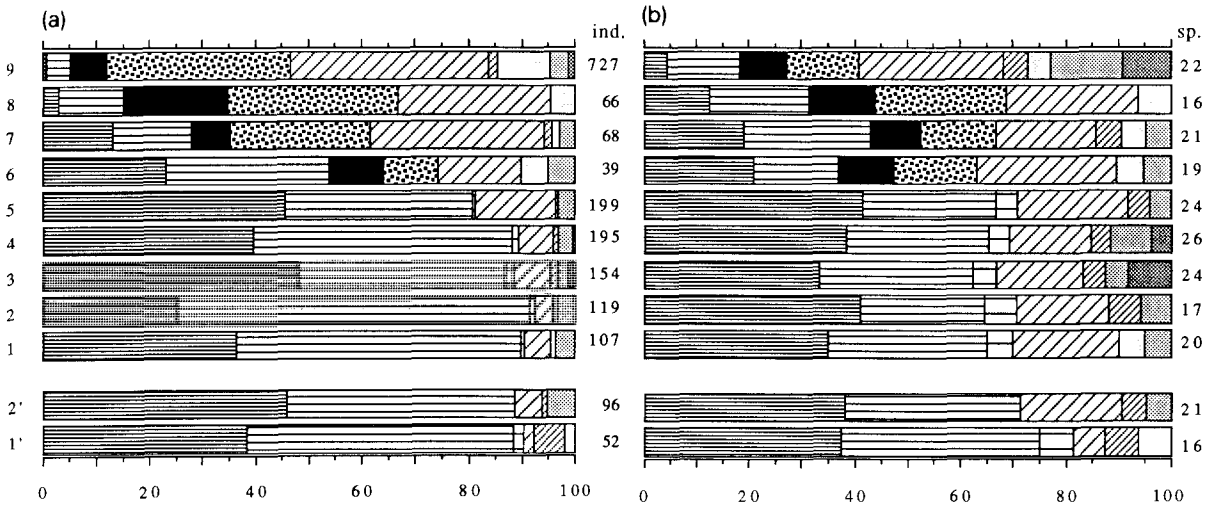


Fig. 7. Vernon. New malacological investigations. Variation of the percentage in (a) individuals (*ind.*) and (b) species (*sp.*) within each ecological group, according to Puisségur (1976) (see legends in Fig. 4).

senting the degradation phase. Under these conditions the correlation with the previous samples (Table 1), a few meters away, is proposed as follows.

1972	new study
Sample 5	sample 6
Sample 4	sample 5
	sample 4
Sample 3	sample 3
Sample 2	sample 2
Sample 1	sample 1

The travertine samples were dated using the <sup>230</sup>Th/<sup>234</sup>U method. The measurements indicate a radioactive equilibrium implying an age equal to or older than 350,000–400,000 yr B.P. (Lécolle et al., 1990). We suggest a correlation with isotopic stage 11, in agreement with the geomorphological and stratigraphical context of this area (Lécolle, 1989).

The composition of the malacofaunas, provided by the two series of samples in Vernon, confirms the biostratigraphic correlation proposed between Vernon and La Celle-sous-Moret (Bourdier et al., 1969). In these two sites, the tufaceous sedimentation is connected with a biological association which gathers both plant and molluscs associa-

tions. According to Clement's (1949) definition of a biome, which evokes that animal populations are associated with formations of wide vegetal complexes, a biome being thus a community of plants and animals usually of the rank of a formation which in other words represents a climatic vegetal and animal community, we can determine the *Lyrodiscus* Biome for this particular Pleistocene community (Rousseau, 1992). Its malacological and botanical composition (occurrence of the fig tree, the laurel of the Canary islands, *Buxus*, *Celtis*) is clearly heterogenous compared to the present time.

### St Pierre-lès-Elbeuf

In the Seine Valley, 42 km west of Vernon (Fig. 1), the St Pierre-lès-Elbeuf deposits yield an important continental record of the Middle and Upper Pleistocene (Lautridou et al., 1974). This locality is a key loess section for Western Europe (Lautridou, 1985). Five non-calcareous loess units were recognized, separated by four interglacial Bt soils: Elbeuf I = Eemian, Elbeuf II, III = intra-Saalian, and Elbeuf IV (more weathered than the others) of Holsteinian age (Lautridou et al., 1974). A calcareous tufa, which yielded a rich malacofauna, overlays this soil (Fig. 8).



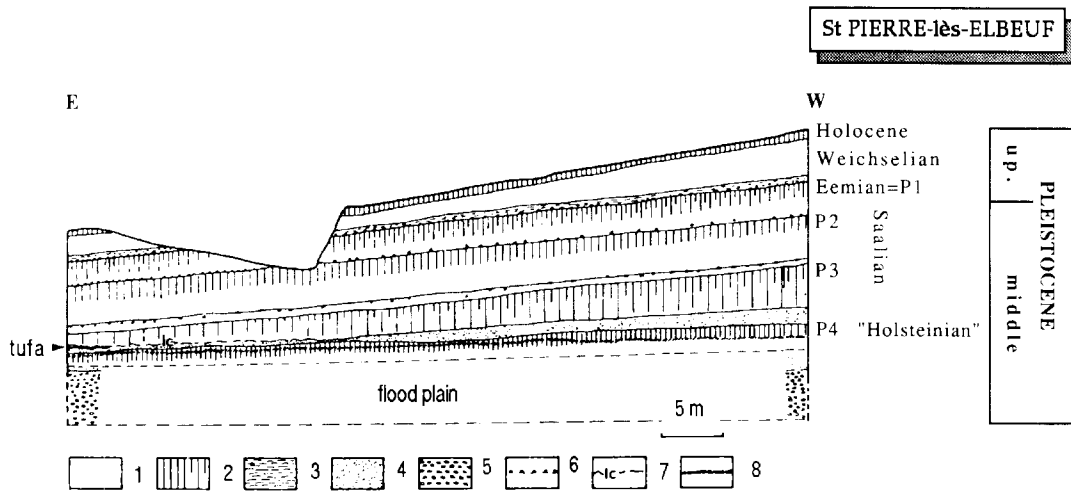


Fig. 8. St. Pierre-lès-Elbeuf. Main outcrop showing the stratigraphic sequence. 1 = loess, 2 = paleosol, 3 = humiferous soils, 4 = fluviatile sands, 5 = fluviatile gravels, 6 = stone line, 7 = "liseré chocolat", 8 = base of the quarry. The "liseré chocolat" corresponds to a thin brown layer of argilans (paleosol III) fixed upon the calcareous tufa (Bt) (from Lautridou et al., 1983, modified).

The malacofauna is typically interglacial (Table 3). Thirteen forest species correspond to 33.8% and 9 semi-forest species to 32.3% of the total amount. The occurrence of species representative of steppe, open ground, xerophilous ecological groups seems to be accidental. Aquatic species are well represented, but the St Pierre malacofauna cannot be an aquatic dominant fauna since it typically characterizes a moist forest where intermittent streams supply semi-permanent pools (Fig. 9a: A, B).

Fossil species, *Retinella (Lyrodiscus) skertchlyi*, *Aegopinella bourdieri*, as well as *Laminifera pauli* (endemic in the Bayonne neighborhood) and *Hygromia cinctella* (Fig. 3), a mediterranean taxon (Table 3) occur as well.

The proximity of both Vernon and St Pierre-lès-Elbeuf permits a correlation to be made based on geomorphological (Lécolle, 1989) and biostratigraphical evidence. Holsteinian age is suggested

for the Vernon upper tufa, in agreement with the radiometric age reported by Lécolle et al., 1990.

*Arrest*

The Arrest site is located in a small valley of the Avalasse river, 8 km south of the Somme estuary (Fig. 1). Tufas occur 22 m above the present-day river level (Fig. 10). These tufas impregnated the underlying chalk which is consolidated. The following succession was observed from the base to the top (Bourdier, pers. comm.): a silty tufa with white concretions providing malacofaunas, a non-cemented level, an indurated level in which one meter thick pedological pockets are present, and finally a colluvium layer. A level of rounded flint gravel occurs between the indurated chalk and the tufas.

As determined in the previous sites, the malaco-

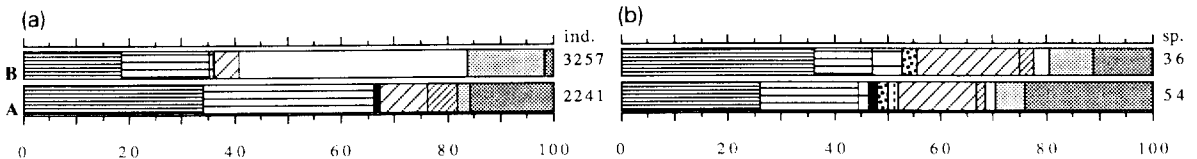


Fig. 9. Variation of the percentage in (a) individuals (ind.) and (b) species (sp.) within each ecological group, according to Puisségur (1976). A = St Pierre-lès-Elbeuf. B = Arrest (see legends in Fig. 4).

TABLE 3

St. Pierre-lès-Elbeuf. List of different species recognized in the samples and individuals counted.

Species		/ Ecol. Groups
<i>Ena montana</i>	25	
<i>Acicula polita</i>	5	
<i>Acanthinula aculeata</i>	57	
<i>Helicodonta obvoluta</i>	60	
<i>Aegopinella pura</i>	212	
<i>Aegopinella nitidula</i>	192	
<i>Cochlodina laminata</i>	35	
<i>Macrogastra rolphii</i>	2	1
<i>Clausilia bidentata</i>	100	
<i>Macrogastra plicatula</i>	4	
<i>Perforatella incarnata</i>	2	
<i>Aegopinella bourdieri</i>	52	
<i>Aegopis acieformis</i>	16	
<i>R. (Lyrodiscus) skertchlyi</i>	1	
<i>Discus rotundatus</i>	345	
<i>Acicula dupuyi</i>	17	
<i>Vitrea crystallina</i>	63	
<i>Laciniaria biplicata</i>	25	
<i>Laminifera pauli</i>	11	
<i>Cepaea hortensis</i>	200	2
<i>Arianta arbustorum</i>	20	
<i>Bradybaena chouquetiana</i>	1	
<i>Bradybaena fruticum</i>	19	
<i>Hygromia cinctella</i>	16	
<i>Macrogastra ventricosa</i>	12	3
<i>Cecilioides acicula</i>	1	4
<i>Vallonia costata</i>	3	5
<i>Monacha cartusiana</i>	2	6
<i>Punctum pygmaeum</i>	25	
<i>Trichia hispida</i>	28	
<i>Vitrina pellucida</i>	5	
<i>Helicigona lapicida</i>	10	
<i>Vitrea contracta</i>	35	7
<i>Euconulus fulvus</i>	1	
<i>Clausilia parvula</i>	93	
<i>Oxychilus sp.</i>	16	
<i>Limax sp.</i>	121	7'
<i>Carychium tridentatum</i>	2	8
<i>Vertigo antivertigo</i>	10	
<i>Succinea putris</i>	27	9
<i>Zonitoides nitidus</i>	17	
<i>Pisidium sp.</i>	22	
<i>Valvata piscinalis</i>	172	
<i>Ancylus fluviatilis</i>	15	
<i>Bathyomphalus contortus</i>	2	
<i>Galba truncatula</i>	5	
<i>Galba palustris</i>	33	
<i>Bulimus sp.</i>	2	10
<i>Radix ovata</i>	9	
<i>Limnaea stagnalis</i>	6	
<i>Gyraulus albus</i>	60	
<i>Planorbis pl. -carinatus</i>	18	
<i>Physa fontinalis</i>	1	
<i>Belgrandia sp.</i>	8	

TABLE 4

Arrest. List of different species recognized in the samples and individuals counted.

Species		/ Ecol. Groups
<i>Acanthinula aculeata</i>	70	
<i>Aegopinella nitidula</i>	176	
<i>Aegopinella pura</i>	149	
<i>Azeca goodalli</i>	50	
<i>Clausilia bidentata</i>	12	
<i>Cochlodina orthostoma</i>	2	
<i>Clausilia sp.</i>	58	1
<i>Discus ruderatus</i>	57	
<i>Ena montana</i>	9	
<i>Helicodonta obvoluta</i>	4	
<i>R. (Lyrodiscus) skertchlyi</i>	11	
<i>Vertigo pusilla</i>	1	
<i>Vitrea subrimata</i>	7	
<i>Arianta + Bradybaena</i>	7	
<i>Cepaea sp.</i>	30	
<i>Discus rotundatus</i>	224	2
<i>Vitrea crystallina</i>	284	
<i>Perforatella bidentata</i>	15	
<i>Macrogastra ventricosa</i>	2	3
<i>Vallonia costata</i>	8	5
<i>Clausilia dubia</i>	15	
<i>Cochlicopa lubrica</i>	14	
<i>Euconulus fulvus</i>	5	
<i>Helicigona lapicida</i>	1	7
<i>Nesovitrea hammonis</i>	3	
<i>Trichia hispida</i>	80	
<i>Vitrea contracta</i>	36	
<i>Limax sp.</i>	1	7'
<i>Carychium tridentatum</i>	1399	8
<i>Carychium minimum</i>	466	
<i>Succinea putris</i>	4	9
<i>Vertigo antivertigo</i>	2	
<i>Anisus leucostomus</i>	24	
<i>Galba palustris</i>	4	
<i>Galba truncatula</i>	5	10
<i>Pisidium sp.</i>	22	

fauna is typically interglacial (Table 4). Forest and semi-forest species represent 52.78% of the total number of species and 35.86% of the total number of individuals (Fig. 9b: A, B). The environment is the same as that of St Pierre-lès-Elbeuf. A moister forest accompanied the occurrence of *Perforatella bidentata*, *Macrogastra ventricosa* and a high number of *Carychium tridentatum* (Table 4). Contrary to the two previous sites, few allochthonous species occur. Only the fossil *Retinella (Lyrodiscus)*

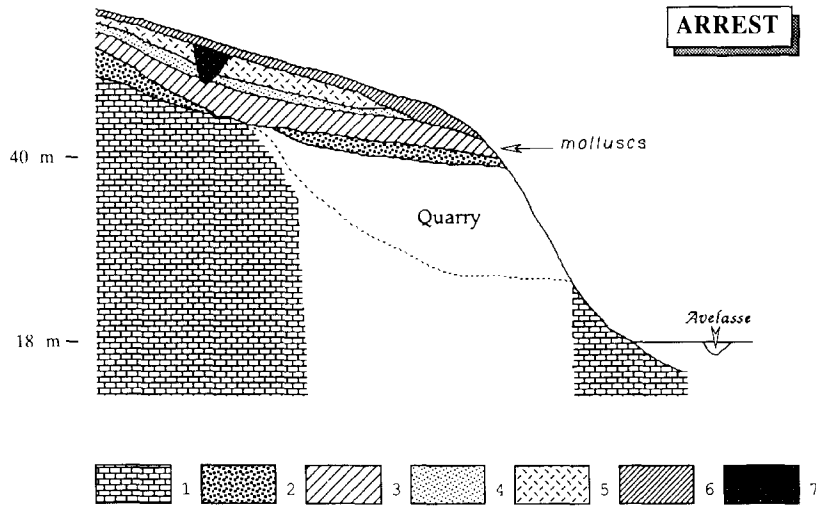


Fig. 10. Arrest. Schematic distribution of the stratigraphy. 1 = substratum (chalk); 2 = rounded flint gravels; 3 = white silty tufa with some concretions; 4 = non-cemented tufa; 5 = indurated tufa; 6 = colluvium; 7 = pedological pocket (from Bourdier, pers. comm., modified)

*cus) skertchlyi* and the central European *Discus ruderatus* were recognized.

The occurrence of these deposits in the Somme Valley implies that this particular assemblage is not a response to the microclimate of the Seine Valley. At least four localities, in an area of about 210 km long, recorded this ecological event. Similar assemblages in the surrounding countries of this area, already published, are examined to determine whether further evidence of this occurrence exist.

**Other West-European malacological records**

*England*

Kerney described a particular malacofauna from tufas in Icklingham (1976) and Hitchin (Kerney, 1959; Holyoak et al., 1983). These two assemblages, typifying an interglacial forest environment (fig. 11a: a, b), are characterized by the occurrence

of species which have no relatives today in the British Isles (Table 5). The assemblages contain the fossil *Acicula diluviana* and *Retinella (Lyrodiscus) skertchlyi*, the central European *Semilimax semilimax*, *Clausilia parvula* and *Acicula polita*, the East European *Ruthenica filograna* and the West European *Laminifera pauli* (Fig. 3). The age allocated to these tufas is Hoxnian, the British equivalent of the Holsteinian (Kerney, 1977).

*Germany*

Münzig (1974) described the malacofaunas from a carbonaceous silty level in the Swabian Jura, at Hörllis, southwestern Germany (Table 6). Once more the composition is heterogeneous and seems to be similar to those already found in France and England (Fig. 11b: A, B). Apart from the extinct *R. (Lyrodiscus) skertchlyi*, the Atlantic species *Clausilia bidentata*, and *Azeca goodalli* occur

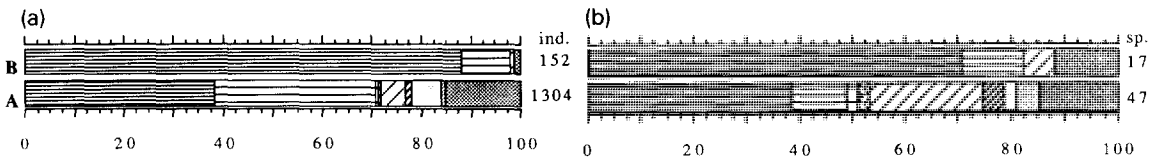


Fig. 11. Variation of the percentage in (a) individuals (*ind.*) and (b) species (*sp.*) within each ecological group, according to Puisségur (1976). A = Icklingham (England). B = Hörllis (Germany) (see legends in Fig. 4).

TABLE 5

Icklingham. List of different species recognized in the samples and individuals counted (from Kerney, 1976).

Species		/ Ecol. Groups
Acanthinula aculeata	50	
Acicula diluviana	5	
Acicula polita	64	
Aegopinella nitidula	78	
Aegopinella pura	98	
Azeca goodalli	92	
Clausilia bidentata	2	
Cochlodina laminata	1	
Cochlodina sp.	5	
Clausilia sp.	28	1
Ena montana	2	
Helicodonta obvoluta	4	
R. (Lyrodiscus) skertchlyi	37	
Ruthenica filograna	1	
Vertigo pusilla	6	
Vitrea subrimata	7	
Zenobiella subrufescens	1	
Zonitoides excavatus	20	
Arianta arbustorum	3	
Cepaea nemoralis	3	
Discus rotundatus	289	2
Pomatias elegans	25	
Vitrea crystallina	101	
Clausilia pumila	1	3
Vallonia costata	4	5
Balea perversa	1	
Clausilia parvula	1	
Cochlicopa sp.	2	
Euconulus fulvus	10	
Helicigona lapicida	2	7
Nesovitrea hammonis	3	
Oxychilus cellarius	3	
Oxychilus sp.	2	
Punctum pygmaeum	30	
Vitrea contracta	17	
Deroceras/Limax	12	
Milax sp.	7	7'
Carychium tridentatum	78	8
Carychium minimum	7	
Leiostyla anglica	3	9
Bathymophalus contortus	1	
Bythinia sp.	1	
Hippeutis complanatus	2	
Pisidium obtusale	7	10
Planorbis planorbis	2	
Sphacrium corneum	1	
Valvata cristata	175	

Alpine *Isognostoma isognostoma*, *Eucoeresia diaphana*, *Trichia edentula*. At Osterbuch, southern Germany, the same association is described from

TABLE 6

Hörlis. List of different species recognized in the samples and individuals counted (from Münzing, 1974).

Species		/ Ecol. Groups
Acanthinula aculeata	1	
Aegopinella sp.	9	
Azeca goodalli	82	
Clausilia bidentata	7	
Discus perspectivus	11	
Ena montana	1	
Isognostoma isognostoma	6	1
Orcula sp.	8	
R. (Lyrodiscus) skertchlyi	5	
Trichia edentula	1	
Vitrea subrimata	2	
Vitrea sp.	1	
Arianta arbustorum	1	
Eucoeresia diaphana	13	2
Clausilia dubia	1	7
Lymnea sp.	1	
Physa fontinalis	1	10

a silty level between two Alpine tills (Münzing and Aktas, 1987). The fauna is quite similar to the Hörlis assemblage showing that the distribution of the *Lyrodiscus* Biome reached as far as 11°E.

### Discussion and conclusion

As determined from the multidisciplinary investigation of the Biache St Vaast site, a well-defined hierarchy seems to exist between pollen and mollusc records (Sommé et al., 1989). When both remains are in complete agreement, the event can be characterized generally as an interglacial phase. On the contrary, if disagreements occur, mainly in the amplitude of the implied climatic variations, an interstadial episode is implied.

Is this warmest stage, recorded by molluscs, also recognized in the pollen assemblages? First, some fossil leaves and seeds indicate the occurrence of an interglacial warm forest in Vernon and La Celle-sous-Moret (Bourdier et al., 1969). Secondly, based on the continental equivalent of isotopic stage 11, several pollen analyses indicated that the related stratigraphical levels indicate a warm climate.

In Great Britain, Hoxnian substages III and IV indicate an oceanic climate with forests of mainly evergreen composition. Besides the sparse mixed-

oak forest elements, Lusitanian components (present-day components of the west Irish flora: *Dabaoecia cantabrica*, *Erica mackiana* and *E. ciliaris*), Hiberno-American species (*Eriocaulon septangulare*) or plants which no longer grow there (*Brasenia cf. purpurea*, *Lysimachis punctata*, *Nymphoides cordata* and *Erica scoparia*) occur (Godwin, 1977).

In southwest France, Texier et al. (1983) drew attention to high temperature conditions during the "Mindel-Riss" stage. The flora contains numerous thermophilous taxa (mixed-oak), hygrophilous herbs but also Mediterranean species (*Rhododendron ponticum*, *Erica mediterranea*) and taxa with Lusitanian affinities (*Dabaoecia cantabrica*).

The Tenagghi Philippon sequence (41°10'N, 24°20'E), in Greece, indicates a long and warm episode at the base of the Lekanis complex during which herbs completely disappear (Van der Hammen et al., 1972).

All these continental records imply a warmer climate, deduced from the malacofaunas of the *Lyrodiscus* Biome, which occurred in Europe during the isotopic stage 11. However, other data from both continental and marine records also show evidence of a warmer climate during isotopic stage 11.

In the Hula Basin, in Northern Israel, a long pollen sequence shows a well developed episode, part of the QVI zone, correlated with stage 11. During this stage, steppe elements are highly dominant, corresponding to the interglacial extension (Fuji and Horowitz, 1989).

In the Colombian Bogota-Funza cores (4°N, 74°W), changes of the vegetation are recorded over the last 3.5 myr B.P. (Hooghiemstra, 1989). During stage 11, the altitudinal position of the forest line was higher than 3300 m, quite similar to the current position. The prevailing average temperature of the high plain of Bogota was estimated to be greater than 14.8°C, associated with a low lake level (Hooghiemstra, 1989).

Ruddiman and McIntyre (1976) showed that Atlantic Ocean cores recorded a well-developed interglacial stage 11, with high carbonate content and high <sup>18</sup>O/<sup>16</sup>O values associated with a high sea level. More recently, DSDP 609 core (49°53'N, 24°14'W) in the Northern Atlantic, indicates the

highest carbonate percentages for the last four climatic cycles during stage 11 (Ruddiman et al., 1989). Core V30-97 (41°10'N, 32°56'W) indicates that the largest differences between current and estimated paleo-sea surface temperature for summer and winter (SSTs and SSTw; +5.1°C) also occurred during stage 11 (Ruddiman et al., 1989).

The climatic signal provided by molluscs during stage 11 is in agreement with the results of various pollen analyses which confirm the interglacial characteristics of the *Lyrodiscus* Biome. Moreover, as marine records also indicate warm climatic conditions during this time span, a global warming, in order to microclimatic conditions over Western Europe, can be involved in permitting the development of this particular assemblage. However, this biological association corresponding to isotopic stage 11 never occurred in the following interglacials (i.e. isotope stages 9, 7 and 5). According to the astronomical theory of climate, climatic variations, as a function of insolation, can be explained in terms of changes in the configuration of the three orbital parameters: (1) the longitude of the perihelion, which determines in which season the insolation is maximum, (2) obliquity, which determines where more insolation is directed in summer, and (3) eccentricity, which is a measure of how elliptical the Earth's orbit about the sun is (Berger, 1978, 1982).

Kukla et al. (1981) studied the orbital configuration of each interglacial. They conclude that interglacials have a specific orbital configuration, high obliquity and an autumn perihelion but can present difference in the eccentricity. Stage 11, for example, shows high obliquity, autumn perihelion but low eccentricity. It shares the same characteristics with the Holocene, however. Although stage 11 shows differences with the other interglacials which could involve particular climatic conditions, this does not explain why such a particular assemblage occurred in Western Europe. The variation in the composition of a biome reflects regional ecological impact on the fauna (Rousseau et al., 1990). This explains the differences observed in the malacofaunas of the *Lyrodiscus* Biome between England and Germany, but does not explain, however, where it originated. Because no similar associations were recorded prior to this stage, in France and Eng-

recorded prior to this stage, in France and England, southern European refuge zones of the declined Tertiary biota can eventually be involved, but need to be investigated. This proposal is mainly based on *R.(Lyrodiscus)*, the history of which seems to be closely connected to that one of the Tertiary forest (Rousseau and Puisségur, 1990b), which occurred in Europe and Africa, and declined to be restricted, for one part, to the Canary archipelago (Raven and Axelrod, 1974). A precise location of these refuge zones is, today, difficult to draw, but the Spanish Atlantic coast or North-western Africa can be proposed. In this way, a limitation, climatic or other, in the opportunities to extend their distribution, can explain the lack of the *Lyrodiscus* Biome during the previous stages. On the contrary, the warmer climatic conditions during stage 11 somehow permitted the molluscs and the plants to occur northward of their restrictive zone. Because such conditions never occurred afterwards and because such assemblages evolved in endemic conditions, the *Lyrodiscus* Biome appears to be a biostratigraphical as well as a biogeographical or climatic marker for the Quaternary in western Europe.

### Acknowledgements

We want to thank J.P. Lautridou for his field assistance and discussions, A. Berger for making his data available, G. Kukla and J. Sommé for fruitful discussions, J. Gavin for improving the manuscript by her criticisms and comments. Contribution CNRS-INSU-DBT n° 380.

### References

- Berger, A., 1978. Théorie astronomique des paléoclimats, une nouvelle approche. *Bull. Soc. Belg. Géol.*, 87: 9–25.
- Berger, A., 1982. Numerical values of the elements of the Earth's Orbit from 5 000 000 YBP to 1 000 000 YAP (Astronomical solution of Berger, 1978). *Inst. Astron. Geophys. Catholic Univ. Louvain-la-Neuve Contrib.*, 35.
- Bourdier, F., Chaline, J. and Puisségur, J.J., 1969. Données nouvelles sur les mollusques et les micromammifères quaternaires des régions d'Amiens et de Paris. *C. R. Acad. Sci. Paris*, 268: 266–269.
- Clements, F.E., 1949. *Dynamics of the Vegetation*. Wilson, New York, NY, 296 pp.
- Fuji, N. and Horowitz, A., 1989. Brunhes epoch paleoclimates of Japan and Israel. In: G. Kukla (Editor), *Long Continental Records of Climate*. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 72: 79–88.
- Godwin, H., 1977. Quaternary history of the British flora. In: F.W. Shotton (Editor), *British Quaternary Studies. Recent Advances*. Clarendon, Oxford, pp. 105–118.
- Holyoak, D.T., Ivanovich, M. and Preece, R.C., 1983. Additional fossil and isotopic evidence for the age of the interglacial tufas at Hitchin and Icklingham. *J. Conchol.*, 31: 260–261.
- Hooghiemstra, H., 1989. Quaternary and upper-Pliocene glaciations and forest development in the tropical Andes: evidence from a long high-resolution pollen record from the sedimentary basin of Bogotá, Colombia. In: G. Kukla (Editor), *Long Continental Records of Climate*. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 72: 11–26.
- Imbrie, J., Hays, J.D., Martinson, D.G., McIntyre, A., Mix, A.C., Morley, J.J., Pisias, N.G., Prell, W.L. and Shackleton N.J., 1984. The orbital theory of Pleistocene climate: support from a revised chronology of the Marine  $\delta^{18}\text{O}$  record. In: A. Berger, J. Imbrie, J. Hays, G. Kukla and B. Saltzman (Editors), *Milankovitch and Climate*. Reidel, Dordrecht, 1, pp. 269–305.
- Kerney, M.P., 1959. An interglacial tufa near Hitchin, Hertfordshire. *Proc. Geol. Assoc.*, 70: 322–337.
- Kerney, M.P., 1963. Late-glacial deposits on the Chalk of south-east England. *Philos. Trans. R. Soc. London, B* 246: 203–254.
- Kerney, M.P., 1976. Mollusca from an interglacial tufa in East Anglia, with the description of a new species of *Lyrodiscus* Pilsbry (Gastropoda: Zonitidae). *J. Conchol.*, 29: 47–50.
- Kerney, M.P., 1977. British Quaternary non-marine Mollusca: a brief review. In: F.W. Shotton (Editor), *British Quaternary Studies. Recent Advances*. Clarendon, Oxford, pp. 31–42.
- Kerney, M.P., Cameron, R.A.D and Jungbluth, J.H., 1983. *Die Landschnecken Nord- und Mitteleuropas*. Parey, Hamburg, 384 pp.
- Kukla, G., Berger, A., Lotti, R. and Brown, J., 1981. Orbital signature of interglacials. *Nature*, 290: 295–300.
- Lautridou, J.P., 1985. *Le cycle périglaciaire Pléistocène en Europe du Nord-Ouest et plus particulièrement en Normandie*. Thesis. Univ. Caen, 908 pp.
- Lautridou, J.P., Masson, M., Paepe, R., Puisségur, J.J. and Verron, G., 1974. Loess, nappes et tuf de Saint-Pierre-lès-Elbeuf, près de Rouen; les terrasses de la Seine de Muids à Caude bec. *Bull. Assoc. Fr. Etud. Quat.*, 40–41: 196–201.
- Lécolle, F., 1989. *Le cours moyen de la Seine au Pléistocène moyen et supérieur*. Géologie et Préhistoire. Groupe Seine, Caen, 549 pp.
- Lécolle, F., Rousseau, D.D., Lautridou, J.P. and Puisségur, J.J., 1990. Le tuf de Vernon: nouvelles données (stratigraphie, paléoclimatologie, datations, corrélations). *Actes Coll. Tufts, Rouen, 1989*. *Bull. Cent. Géomorphol. Caen*, 38: 131–149.
- Ložek, V., 1964. Quatärmollusken der Tschechoslowakei. *Rozpr. Ustred. Ustavu. Geol.*, 31: 1–374.
- Münzing, K., 1974. Mollusken aus dem älteren Pleistozän Schwabens. *Jahresh. Geol. Landesamt Baden-Württemberg*, 16: 61–78.
- Münzing, K. and Aktas, A., 1987. Weitere Funde molluskenführender Mergellagen im Unteren Deckshotter von Bayer-

- isch Scwaben. Jahrber. Mitt. Oberrhein. Geol. Ver. N.F., 69: 181–193.
- Peteet, D.M., Vogel, J.S., Nelson, D.E., Southon, J.R., Nickmann, R.J. and Heusser, L.E., 1990. Younger Dryas climatic reversal in northeastern USA ? AMS ages for an old problem. *Quat. Res.*, 33: 219–230.
- Puisségur, J.J., 1976. Mollusques continentaux quaternaires de Bourgogne. Significations stratigraphiques et climatiques. Rapports avec d'autres faunes boréales de France. *Mém. Géol. Univ. Dijon*, 3, 241 pp.
- Raven, P.H. and Axelrod, D.I., 1974. Angiosperm biogeography and past continental movements. *Ann. Mo. Bot. Garden*, 61: 539–673.
- Roussignol-Strick, M., 1983. African monsoons, an immediate climate response to orbital insolation. *Nature*, 303: 46–49.
- Rousseau, D.D., 1992. Is causal ecological biogeography a progressive research program ? *Quat. Sci. Rev.*, in press.
- Rousseau, D.D. and Puisségur, J.J., 1989. *Aegopinella bourdieri* nov. sp. (Gastropoda: Zonitidae), une espèce nouvelle du Pléistocène normand. *Géobios*, 22: 123–128.
- Rousseau, D.D. and Puisségur, J.J., 1990a. A 350,000 years climatic record from the loess sequence of Achenheim, Alsace, France. *Boreas*, 19: 203–216.
- Rousseau, D.D. and Puisségur, J.J., 1990b. Phylogénèse et biogéographie de *Retinella* (*Lyrodiscus*) *Pilsbry* (Gastropoda; Zonitidae). *Géobios*, 23: 57–70.
- Rousseau, D.D., Puisségur, J.J. and Lautridou, J.P., 1990. Biogeography of the Pleistocene pleniglacial malacofaunas in Europe. Stratigraphic and climatic implications. In: D.D. Rousseau (Editor), *Methods and Concepts in European Quaternary Stratigraphy*. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 80: 7–23.
- Ruddiman, W.F. and McIntyre, A., 1976. Northeast Atlantic paleoclimatic changes over the past 600,000 years. *Geol. Soc. Am. Bull.*, 145: 111–146.
- Ruddiman, W.F., Raymo, M.E., Martinson, D.G., Clement, B.M. and Backman, J., 1989. Pleistocene evolution: northern hemisphere ice sheets and North Atlantic Ocean. *Paleoceanography*, 4: 353–412.
- Sommé, J., Tuffreau, A., Aitken, M., Auguste, P., Chaline, J., Colbeaux, J.P., Hus, J., Huxtable, J., Juvigné, E., Munaut, A.V., Occhiotti, S., Puisségur, J.J., Rousseau, D.D. and Van Vliet-Lanoë, B., 1989. Chronostratigraphie, climats et environnements. In: A. Tuffreau et J. Sommé (Editors), *Biache-St Vaast, I. Mém. Soc. Préhist. Fr.*, 21: 115–119.
- Texier, J.P., Raynal, J.P., Laville, H., Paquereau, M.M., Prat, F., Debenath, A. and Delpech, F., 1983. Histoire climatique de l'Aquitaine du Pléistocène ancien au dernier interglaciaire. *Bull. Inst. Géol. Bassin Aquitaine*, 34: 207–217.
- Van der Hammen, T., Wijnstra, T.A. and Zagwijn, W.H., 1972. The floral record of the Late Cenozoic of Europe. In: K.K. Turekian (Editor), *The Late Cenozoic Glacial Ages*. Yale Univ. Press, New Haven, CN, pp. 391–424.
- Woillard, G., 1978. Grande Pile peat bog: a continuous pollen record for the last 140,000 years. *Quat. Res.*, 9: 1–21.