

Long distance transport of pollen to Greenland

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[1] Four pollen traps were recently placed at coastal sites in East and West Greenland in order to assess long distance transport of pollen to the Arctic domain. By identifying potential vegetation source areas associated with air mass pathways we were able to produce the first detailed record of pollen transport from eastern North America to Narsarsuaq, southern Greenland. The record is based on observations and on a transport and dispersion model used to calculate back trajectories of the air masses. The evidence points to the pollen being transported northward by the air mass over Newfoundland and the Labrador Sea before reaching Narsarsuaq. At this point, the air mass was at an altitude of 3000 m. Deposition of pollen grains occurred with downward air movements associated with a 0.5 mm/h rain. The source areas for these pollen grains differ from those of dust and certain other aerosols that reach the summit of the Greenland Ice Sheet. The results demonstrate the need for continued maintenance and analysis of the Greenland pollen trap data in order to improve our understanding of atmospheric circulation and transport to high northern latitudes.

INDEX TERMS: 0315 Atmospheric Composition and Structure: Biosphere/atmosphere interactions; 0305 Atmospheric Composition and Structure: Aerosols and particles (0345, 4801); 3322 Meteorology and Atmospheric Dynamics: Land/atmosphere interactions; 3349 Meteorology and Atmospheric Dynamics: Polar meteorology; 1620 Global Change: Climate dynamics (3309). **Citation:** Rousseau, D.-D., D. Duzer, G. Cambon, D. Jolly, U. Poulsen, J. Ferrier, P. Schevin, and R. Gros, Long distance transport of pollen to Greenland, *Geophys. Res. Lett.*, 30(14), 1765, doi:10.1029/2003GL017539, 2003.

1. Introduction

[2] Although dust and other aerosol tracers have been used to infer both present and past atmospheric circulation patterns [Biscaye *et al.*, 1997; Bory *et al.*, 2002; Goto-Azuma and Koerner, 2001; Kahl *et al.*, 1997], other attempts have been made to use pollen as an air mass tracer. Studying the modern pollen deposition in the Canadian Arctic and subarctic, Ritchie and Lichti-Federovich [1967] indicate the occurrence of long distance pollen transport of anemophilous trees growing in the temperate North American zone, of Canada and Alaska. Andrews *et al.* [1980], Bourgeois *et al.* [1985], Gajewski [1995], Jacobs *et al.* [1985], Nichols [1967], Ritchie [1974], Ritchie *et al.* [1987],

and Short and Holdsworth [1985] show different transports according to the distance traveled from local, extralocal, regional to extraregional as defined by Janssen [1973]. However, most of these examples correspond to transport of not more than about 500 km. More recently, while sampling snow and ice in the Canadian Arctic, Bourgeois [Bourgeois, 2000; Bourgeois *et al.*, 2001] identified the occurrence of seasonal long distance pollen transport along NS transects starting from the Canadian boreal forest zone to the North Pole. In Europe, pollen transport to the Scandinavian Arctic was characterized through an annual survey of pollen sedimentation [Hicks *et al.*, 2001].

[3] Most of these studies lack a regular and quantified survey of the pollen dispersion. With the exception of rare events such as those reaching Northern Europe [Franzen *et al.*, 1994; Hjelmroos and Franzen, 1994], or the High Canadian Arctic [Campbell *et al.*, 1999], pollen grains marking long distance transport are generally in very low number. The analysis by Bourgeois *et al.* [2001] of the Arctic Ocean ice supports this interpretation as she determined pollen grains from oak, elm, maple, ash, typical of deciduous forest south of the North American boreal forest but in very low concentration.

[4] A network of four pollen traps have been established at coastal sites in East and West Greenland during the last two years as part of the (EPILOBE Project [Rousseau *et al.*, 2001]) (Figure 1). We present here results for Narsarsuaq (61.15° N, 45.43° W, 1 m asl) (Figure 2A) in southern Greenland whose modern soil samples have yielded allochthonous pollen taxa such as *Pinus* (pine), *Picea* (spruce), *Quercus* (oak) and *Olea* (olive). Greenland is very sparsely

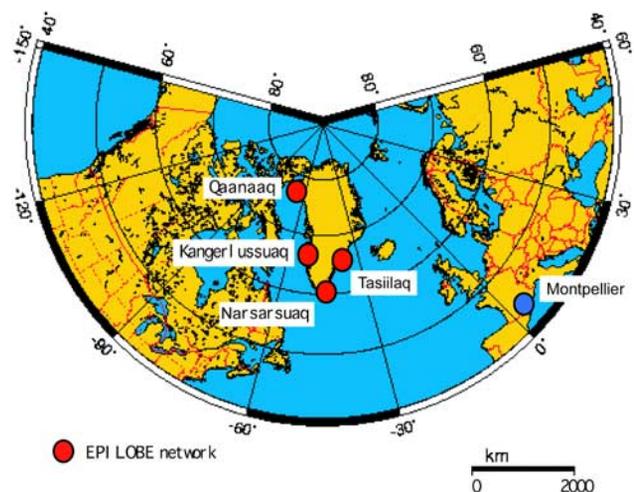


Figure 1. Location the EPILOBE network pollen trap sites, Greenland.

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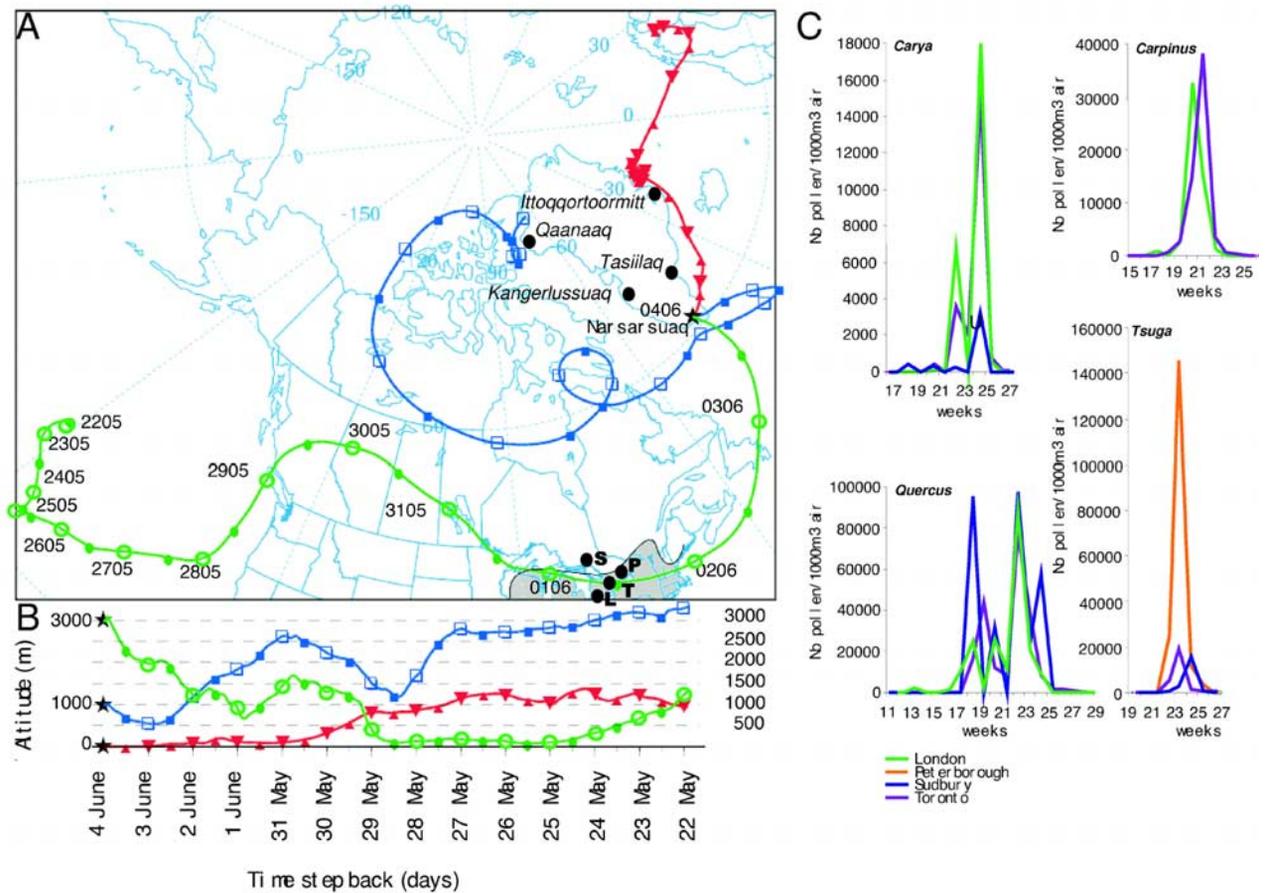


Figure 2. Long distance transport to Greenland. (a) Backward trajectories provided by the HYSPLIT model [HYSPLIT4Model, 1997] of air masses reaching Narsarsuaq (61.15° N, 45.43° W, 1 m asl) at different altitudes: ground level (red), 1000 m (blue) and 3000 m (green) on June 4th of 2002. The grey area in eastern USA and south-eastern Canada represents the zone where *Quercus* (oak), *Fagus* (beech), *Carpinus* (hornbeam), *Tsuga* (hemlock) and *Carya* (hickory) are all growing [from Thompson et al., 1999a; 1999b]. The “3000” m air mass passed over this area. L: London, P: Peterborough, S: Sudbury, T: Toronto. (b) Altitude variation of the three air masses used in the backward trajectories analysis. The “3000 m” air mass over Narsarsuaq on June 4th, 2002, was at a lower elevation on June 1st, when it passed over the area where *Quercus*, *Fagus*, *Carpinus*, *Tsuga* and *Carya* are growing. (c) Number of pollen grains captured per thousand of cubic meters of air versus weeks expressed by their yearly occurrence number at four pollen stations located in southern Ontario, Canada (London 43°N, 81.25°W; Peterborough 44.33°N, 78.33°W; Toronto 43.65°N, 79.33°W; Sudbury 46.50°N, 81°W) for *Carya*, *Carpinus*, *Quercus* and *Tsuga*. Week 17 started on April 23, 18 on April 30, 19 on May 7, 20 on May 14, 21 on May 21, 22 on May 28, 23 on June 4, 24 on June 11 and 25 on June 18. [from Cambon, 1994; Cambon et al., 1992].

vegetated [Fredskild, 1983; Pennington, 1980] and has a low pollen productivity compared to North American [Cambon, 1994; Cambon et al., 1992] and European [Cambon, 1981] regions which simplifies the identification of long distance input. In Narsarsuaq, the local vegetation is composed of shrubs and dwarf shrub of *Betula* (birch), *Salix* (willow), *Juniperus* (juniper), with associated grasses and sedges, typical of low and mid arctic tundra.

2. Results and Discussion

2.1. Pollen Occurrence

[5] Our pollen traps [Cour, 1974], unlike the more commonly used Tauber trap, consist of sticky filters mounted on a weather vane, which allows them to be constantly exposed to the wind. The filters are replaced

every two weeks at this southern site. In addition to the local vegetation, five taxa, allochthonous to Greenland, have been identified for the period June 3–16, 2002 (23rd and 24th weeks). This is the only time since the installation of the pollen trap, in June 2001, that pollen grains, not indigenous to Greenland, have been found. Among the 5514 counted pollen grains, three are identified as *Quercus* (oak), two as *Fagus* (beech), two as *Carya* (hickory), one as *Tsuga* (hemlock) and two as *Carpinus* (hornbeam) and one as *Juglans* (walnut). Eight grains of *Pinus* (pine) were also found. However, because *Pinus* is so widely distributed and easily transported from different northern areas, we did not include it in our analysis. The shape of the transported pollen grains is severely altered indicating long distance (>1000 km). A replicate analysis of the sample provides similar results.

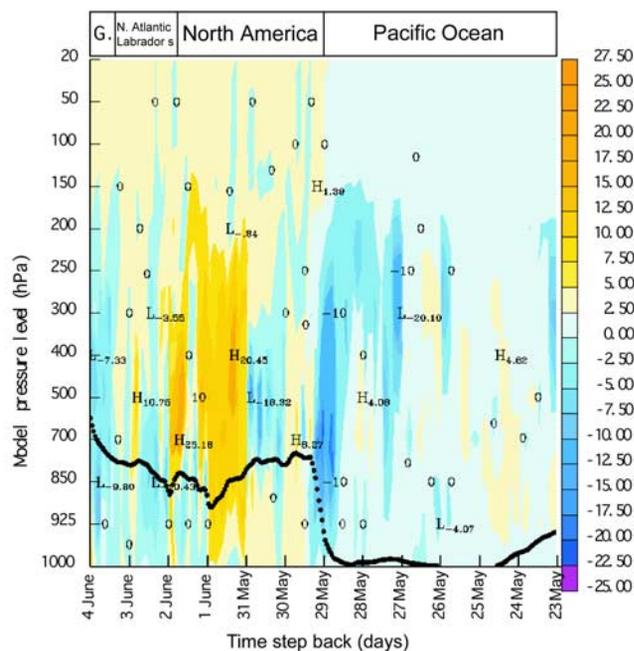


Figure 3. Updrafts and downdrafts in the atmosphere. Velocity of the air mass passing over *Quercus* (oak), *Fagus* (beech), *Carpinus* (hornbeam), *Tsuga* (hemlock) and *Carya* (hickory) growing area in North America and reached Narsarsuaq at 3000 m above the station on June 4, 2002 versus time (see Figure 2 for the geographic trajectory). Yellow to brown values indicate upward movements whereas light blue to pink indicate downward movements. The dark line corresponds to the air mass.

[6] The simultaneous occurrence of *Carya* and *Tsuga* in the same pollen sample suggests that eastern North America is the most probable source of these allochthonous pollen grains. These trees do not grow in Europe but are found in the eastern USA and south eastern Canada [Thompson *et al.*, 1999a; Thompson *et al.*, 1999b] (Figure 2a). Moreover, the pollen emission of oak, beech, hickory, hemlock, hornbeam and walnut pollen grains has been observed in the Ontario region of Canada (in London, Peterborough, Sudbury and Toronto) in 1983 and 1984 by Cambon [Cambon, 1994; Cambon *et al.*, 1992] using the same pollen trap system (Figure 2c). The mean monthly temperatures is not significantly different in 2002. The pollen emission data from the four Canadian sampling sites [Cambon, 1994; Cambon *et al.*, 1992], are then used to identify the time interval when pollen is emitted to the atmosphere (Figure 2c). It is expected that this region is also representative of the northeastern USA.

2.2. Backward Trajectories

[7] The HYSPLIT (Hybrid Single-Particle Lagrangian Integrated Trajectory) transport and dispersion model computes trajectories for any place in the world from archives of gridded meteorological data [HYSPLIT4 Model, 1997]. HYSPLIT backward trajectories from Narsarsuaq show that an air mass passed over the northern boundary of the range distribution of the identified trees in the eastern USA on June 1st of 2002 (Figures 2a and 2b). The pollen grains would have been carried upward to about the 895 hPa level

(about 630 m above ground level) (Figure 3), and then transported eastward over the USA and then northward over Newfoundland and the Labrador Sea to Greenland. They reached Narsarsuaq early in the morning of June 4th with the air mass at an altitude of 3000 m (Figure 2). The pollen grains were deposited with downward air movements associated with a 0.5 mm/h rain in the early morning (Figure 3). The analysis of the 3D backward trajectories shows that several previous downward air motions are likely to have contributed to considerably reduce the number of pollen grains being deposited at this site. The model indicates that just when the air mass passed over the growing area of oak, beech, hornbeam, hemlock and hickory, upward air movements could have lifted the pollen grains to a high altitude during a very limited interval, but long enough to begin their long distance transport.

[8] The identified transport from eastern northeastern North America to southern Greenland took three days. Barry *et al.* [1981] reported a 5 days average duration for a pollen grain to be transported, which is similar to the transport we identified. Comparing these pollen results with those of dust transport to Greenland, it appears that the source of the air masses are not the same. Dust in the ice at the summit of the Greenland ice cap has been shown to have originated mostly from the Takla Makan and Gobi deserts [Biscaye *et al.*, 1997; Bory *et al.*, 2002]. However, the atmospheric dynamics for both cases are not the same. The general range in size of pollen grains ($10 \mu < \text{Size} < 200 \mu$) place them among the mechanically produced super-micrometer aerosol particles [Charlson and Heintzenberg, 1995]. The pollen grains identified at our site range between 20 and 80 μm . Pollen grains of similar size have been identified in snow deposited on the Greenland Ice Sheet and Canadian Arctic ice caps [Bourgeois, 1990; Bourgeois *et al.*, 2000; Short and Holdsworth, 1985]. Dust on the other hand ranges in smaller size and originated from areas without any vegetation.

[9] Similar to dust or chemical elements, long distance transport of pollen grains requires several conditions. First, the source area must be favorable for both the emission in the air and uplift to a sufficient altitude to be transported. Second, the pollen must remain in the air, and not sink via dry or wet deposition in route towards the sampling station. However, unlike dust, long distance pollen transport must coincide with the timing of the emission of pollen grains in the air. Our case study documents the first long distance pollen transport, from emission to capture. Ongoing analysis is expected to improve understanding of past and present atmospheric circulation and long distance transport of super-micrometer aerosol particles to the high northern latitudes.

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