

Atelier “Analyse des données”

Lecture 2: Singular Spectrum Analysis (SSA) and the SSA-MTM Toolkit



Michael Ghil
CERES-ERTI, etc.



Based on joint work with many students, post-docs, and colleagues over the years; please see

<http://www.environnement.ens.fr/> and

<http://www.atmos.ucla.edu/tcd/> for further details.

Advanced Spectral Methods, Nonlinear Dynamics, and the Nile River

Michael Ghil

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Motivation

1. Climatic time series have typically broad peaks on top of a continuous, “warm-colored” background → *Method*
2. Connections to nonlinear dynamics → *Theory*
3. Need for stringent statistical significance tests → *Toolkit*
4. Applications to analysis and prediction → *Examples*

Joint work with: M. R. Allen, M. D. Dettinger, K. Ide, N. Jiang, C. L. Keppene, D. Kondrashov, M. Kimoto, M. E. Mann, J. D. Neelin, M. C. Penland, G. Plaut, A. W. Robertson, A. Saunders, D. Sornette, S. Speich, C. M. Strong, C. Taricco, Y.-d. Tian, Y. S. Unal, R. Vautard, & P. Yiou (on 3 continents).

<http://www.atmos.ucla.edu/tcd>

Motivation & Outline

1. **Data sets** in the geosciences are often **short and contain errors**: this is both an obstacle and an incentive.
2. **Phenomena** in the geosciences often have both **regular** (“cycles”) and **irregular** (“noise”) aspects.
3. Different spatial and temporal scales:
one person’s noise is **another person’s signal**.
4. Need both **deterministic** and **stochastic** modeling.
5. **Regularities** include **(quasi-)periodicity** → spectral analysis via “classical” methods (see **SSA-MTM Toolkit**).
6. **Irregularities** include **scaling and (multi-)fractality** → “spectral analysis” via Hurst exponents, dimensions, etc. (see **Multi-Trend Analysis, MTA**)
7. Does some **combination of the two**, + **deterministic** and **stochastic** modeling, provide a **pathway to prediction**?

For details and publications, please visit these two Web sites:

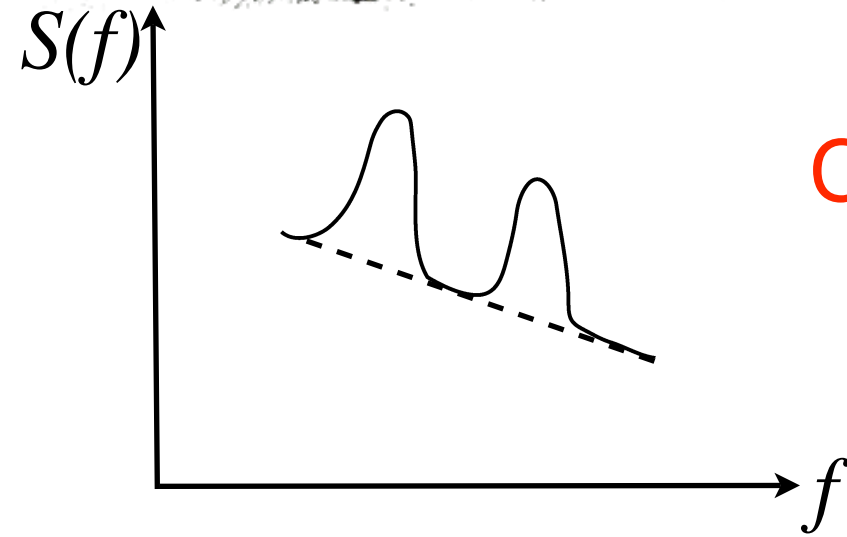
TCD <http://www.atmos.ucla.edu/tcd/> → key person – **Dmitri Kondrashov!**

E2-C2 http://www.ipsl.jussieu.fr/~ypsce/py_E2C2.html

Outline

- Time series analysis
 - The “smooth” and “rough” part of a time series
 - Oscillations and nonlinear dynamics
- Singular spectral analysis (SSA)
 - Principal components in time and space
 - The SSA-MTM Toolkit
- The Nile River floods
 - Longest climate-related, instrumental time series
 - Gap filling in time series
 - NAO and SOI impacts on the Nile River
- Concluding remarks
 - Cautionary remarks (“garde-fous”)
 - References

Spectral Density (Math)/Power Spectrum (Science & Engng.)



Continuous background
+ peaks

◦ Wiener-Khinchin (Bochner) Theorem

Blackman-Tukey Method

$$R(s) = \lim_{L \rightarrow \infty} \frac{1}{2L} \int_{-L}^L x(t)x(t+s)dt$$

$$S(f) = \frac{1}{2\pi} \int_{-\infty}^{\infty} R(s)e^{-ifs}ds \equiv \hat{R}(s)$$

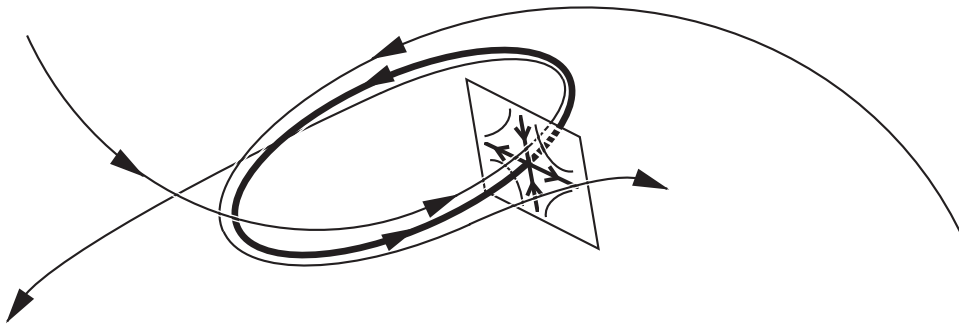
i.e., the lag-autocorrelation function & the spectral density

are Fourier transforms of each other.

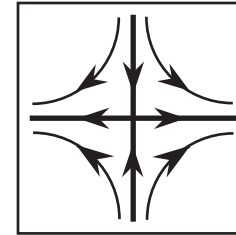
Power Law for Spectrum (cont'd)

- Hypothesis: “**Poles**” correspond to the least unstable periodic orbits

“*unstable limit cycles*”



“*Poincaré section*”



- Major clue to the physics

that underlies the dynamics

- N.B. Limit cycle not necessarily elliptic, i.e. not

$$(x, y) = (a_f \sin(ft), b_f \cos(ft))$$

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Spectral density: deterministic signal

- Let f be a 1-D signal, with possibly complex values, such that $f \in L^1(\mathbb{R}) \cap L^2(\mathbb{R})$. We define the **autocorrelation** of the signal by:

$$R_f(\tau) := \int_{\mathbb{R}} f(x) \overline{f(x - \tau)} dx, \quad \text{for } \tau \in \mathbb{R}. \quad (1)$$

- The **power spectral density** (PSD) of this signal is given by the positive real function:

$$S_f(\xi) = |\hat{f}(\xi)|^2,$$

where $\hat{f}(\xi)$ represents the Fourier transform of f in ξ . The PSD is often called simply the *spectrum* of the signal. Typically it measures the amount of energy per unit of time (i.e. the power) contained in the signal as a function of the frequency ξ .

- If we introduce the notation $f^s(x) := \overline{f(-x)}$ then $R_f = f * f^s$.
- Exercise:** Show that the PSD and the autocorrelation are Fourier transforms of each other, up to a multiplicative constant.

Spectral density: random signal

- Let $f(\omega, \cdot)$ be a **stationary process**, for $\omega \in \Omega$ a probability space. We define the **autocorrelation** of the signal by:

$$R_f(\tau) := \mathbb{E}(f(\omega, t)f(\omega, t + \tau)), \quad (2)$$

where the expectation is taken over all the **realizations** ω . Note that $R_f(\tau)$ is well-defined even if f is not in $L^p(\mathbb{R})$ for any p .

- We assume that $R_f(\cdot)$ is integrable, and that the random process f is **ergodic**. Under these conditions

$$R_f(\tau) = \lim_{T \rightarrow \infty} \frac{1}{T} \int_{-T/2}^{T/2} f(t)f(t + \tau)dt$$

- We define the truncated Fourier transform, $F_T(\xi) := \int_{-T/2}^{T/2} f(t)e^{-2i\pi\xi t}dt$, and the truncated power spectral density, $\frac{1}{T}|F_T(\xi)|^2$.
- Theorem (Wiener-Khinchin):** For all ξ , $\lim_{T \rightarrow \infty} \mathbb{E}(\frac{1}{T}|F_T(\xi)|^2)$ exists and is given by

$$\int_{-\infty}^{\infty} R_f(\tau)e^{-2i\pi\xi t}d\tau; \text{ this limit defines the PSD in this context.}$$

Empirical orthogonal functions of a scalar field (I)

The Karhunen-Loève expansion (I)

- Let $Y(t, \mathbf{x})$ be a **random scalar field** (we omit the dependence on ω), for $\mathbf{x} \in D$ a **spatial domain**. We seek the **Karhunen-Loève expansion** (or EOF expansion) of Y , namely:

$$Y(t, \mathbf{x}) := \sum_k Z_k(t) \varphi_k(\mathbf{x}), \quad (3)$$

where $\int_D \varphi_i(\mathbf{x}) \varphi_j(\mathbf{x}) = \delta_{ij} \lambda_i$, and λ_i is the corresponding eigenvalue.

- The random functions $\{Z_i(t)\}$ are said to be **statistically orthogonal** if $\mathbb{E}(Z_i(t) Z_j(t)) = \delta_{ij} \lambda_i$; they are found by projecting each realization onto the EOFs: $Z_k(t) := \int_D Y(t, \mathbf{x}) \varphi_k(\mathbf{x}) d\mathbf{x}$.
- The eigenfunctions are the solutions of the **Fredholm integral equation**

$$\int_D R(\mathbf{x}, \mathbf{y}) \varphi_k(\mathbf{y}) d\mathbf{y} = \lambda_k \varphi_k(\mathbf{x}),$$

where $R(\mathbf{x}, \mathbf{y})$ is the **autocorrelation** of the process $Y(t, \mathbf{x})$.

- It can be shown that there is a countably infinite number of such eigenfunctions and

$$R(\mathbf{x}, \mathbf{y}) = \sum_k \lambda_k \varphi_k(\mathbf{x}) \varphi_k(\mathbf{y}).$$

The Karhunen-Loève expansion (II)

- It can be shown that the truncated decomposition $Y_K(t, \mathbf{x}) := \sum_{k=1}^{K} Z_k(t) \varphi_k(\mathbf{x})$ minimizes the mean **integrated squared error**:

$$\mathbb{E} \left\{ \int_D [Y(t, \mathbf{x}) - Y_K(t, \mathbf{x})]^2 d\mathbf{x} \right\} := \sum_{k=K+1}^{\infty} \lambda_k.$$

- The spectral decomposition is optimal in the sense that this error is a minimum compared to K terms of any orthogonal systems (Cohen and Jones, 1969).
- The discrete analogue:** Given observations of the random field Y for a given time at M locations, we have to estimate $R(\mathbf{x}, \mathbf{y})$ and then solve the Fredholm eq. The numerical approximation to this equation is:

$$\sum_{j=1}^M R(\mathbf{x}_i, \mathbf{x}_j) \varphi_k(\mathbf{x}_j) \Delta_j = \lambda_k \varphi_k(\mathbf{x}_i), \quad i = 1, \dots, M,$$

where Δ_j is the element of area associated with the observation at \mathbf{x}_j .

The Karhunen-Loève expansion (III)

- If we let

$$\Gamma_{ij} = R(x_i, x_j) \sqrt{\Delta_i \Delta_j},$$
$$\nu_j^k = \varphi_k(x_i) \sqrt{\Delta_i},$$

then it can be shown that

$$\sum_{j=1}^p \Gamma_{ij} \nu_j^k = \lambda_k \nu_i^k;$$

this represents the eigendecomposition of the **covariance matrix** Γ with eigenvalues $\{\lambda_k\}$. The eigenvectors ν_i^k estimate the eigenfunctions $\varphi_k(x_i)$ at the observation points, multiplied by $\sqrt{\Delta_i}$.

- Singular spectrum analysis of time series extends these ideas from random fields to time series.

Singular spectrum analysis

Continuous formulation (I)

- We consider here a zero-mean, continuous, infinite and **ergodic** time series $X(t)$. We introduce the integral operator A , acting on functions $f(t)$ belonging to $H := L^2(-T, T)$ according to:

$$Af(t) := \frac{1}{2T} \int_{-T}^T R_X(t-s)f(s)ds,$$

where $R_X(\tau)$ is the autocorrelation of X at lag τ .

- Since $R_X(\tau)$ is symmetric w.r.t. the origin $\tau = 0$, A is a symmetric non-negative operator. **Hilbert-Schmidt theory** ensures the existence of a countable set of nonnegative eigenvalues and corresponding eigenfunctions (still called EOFs) of

$$A\varphi = \lambda\varphi.$$

- Each segment of X — having length $2T$ and centered at t — is a function in H that may be expanded w.r.t. the EOF basis. For each t and s , such that $|s| \leq T$, provided $\lambda_k > 0$, for all k ,

$$X(t+s) = \sum_{k=1}^{\infty} X_k(t)\varphi_k(s).$$

Continuous formulation (II)

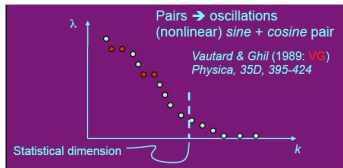
- The projection $X_k(t) := \frac{1}{2T} \int_{-T}^T X(t+s)\varphi_k(s)ds$ onto the k -th EOF constitutes a zero-mean time series which is called the k -th PC. $X_k(t)$ has variance λ_k , and it represents a filtered version of X , the filter being the moving average weighted by φ_k .
- The PCs are **uncorrelated** with each other at zero lag, but their cross-covariances are not vanishing for lags $\neq 0$.
- From Mercer's theorem (Riesz & Nagy, 1955) applied to the Hilbert-Schmidt operator A , it follows that the autocorrelation can be expanded as

$$R_X(t-s) = \sum_{k=1}^{\infty} \lambda_k \varphi_k(t)\varphi_k(s),$$

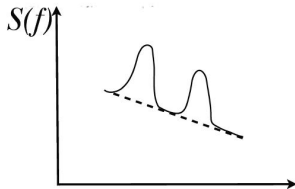
for a.e. t and s in $(-T, T)$. Taking $t = s$ and integrating from $-T$ to T , gives $\sum_{k=1}^{\infty} \lambda_k = R_X(0)$, where $\sigma_X^2 := R_X(0)$ is the **total variance** of X .

"Scree diagram"

Comparison between the "scree diagram" and the power spectrum

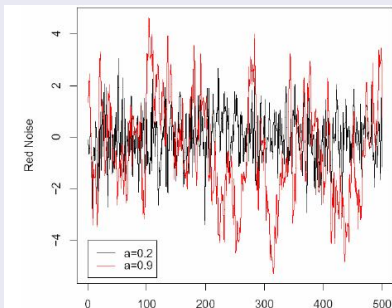


- The equal-eigenvalue **pairs** correspond to the **peaks** in the power spectrum.



"Colored" Noises, I

Red noise



- It is generated from $X(t+1) = aX(t) + b(t)$, with $0 \leq a < 1$ and b is i.i.d. with a Gaussian distribution $\mathcal{N}(0, \sigma)$.

- This process displays more power at low frequencies. Its correlation is

$$R_X(\tau) = \sigma^2 \frac{a^{|\tau|}}{1-a^2}$$

$$S_X(\xi) = R_X(0) \cdot \frac{1-a^2}{1-2a\cos(2\pi\xi) + a^2} \sim \frac{1}{b^2 + c^2\xi^2}$$

Power Spectra & Reconstruction

◦ A. Transform pair:

$$X(t + s) = \sum_{k=1}^M a_k(t) e_k(s), e_k(s) - EOF$$

The e_k 's are **adaptive filters**,

$$a_k(t) = \sum_{s=1}^M X(t + s) e_k(s), a_k(t) - PC$$

the a_k 's are **filtered time series**.

B. Power spectra

$$S_X(f) = \sum_{k=1}^M S_k(f); \quad S_k(f) = R_k(s); \quad R_k(s) \approx \frac{1}{T} \int_0^T a_k(t) a_k(t + s) dt$$

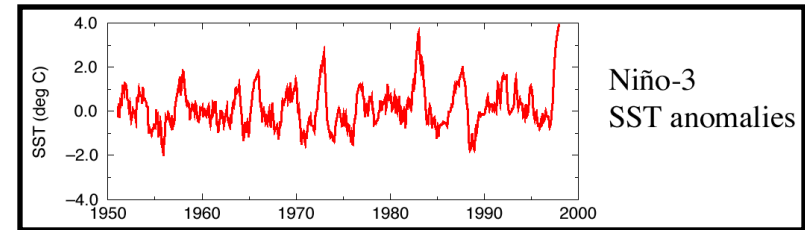
C. Partial reconstruction

$$X^K(t) = \frac{1}{M} \sum_{k \in K} \sum_{s=1}^M a_k(t - s) e_k(s);$$

in particular: $K = \{1, 2, \dots, S\}$ or $K = \{k\}$ or $K = \{l, l + 1; \lambda_l \approx \lambda_{l+1}\}$

Singular Spectrum Analysis (SSA)

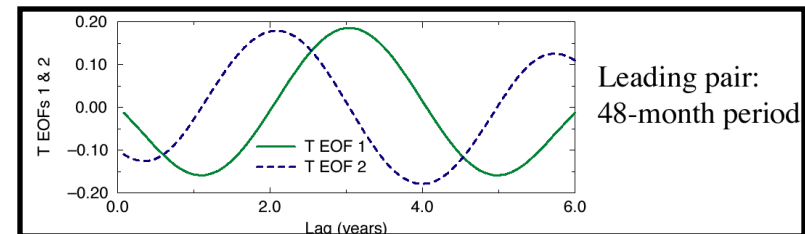
Time series



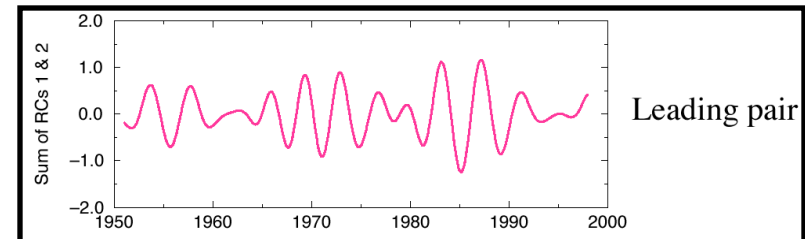
SSA decomposes (geophysical & other)
time series into

Temporal EOFs (T-EOFs) and
Temporal Principal Components (T-PCs),
based on the series' lag-covariance matrix

T-EOFs



RCs

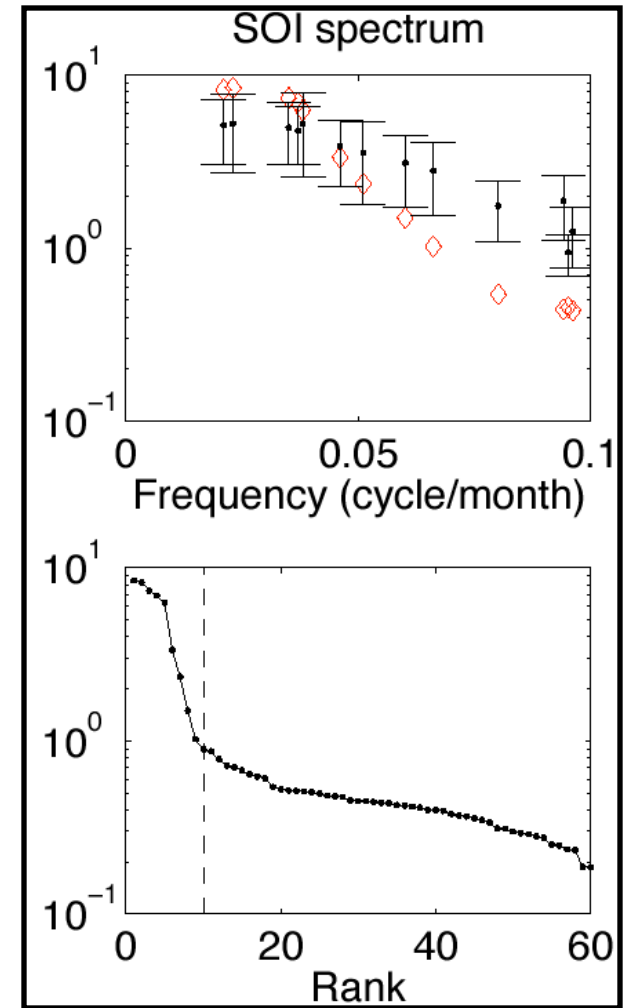
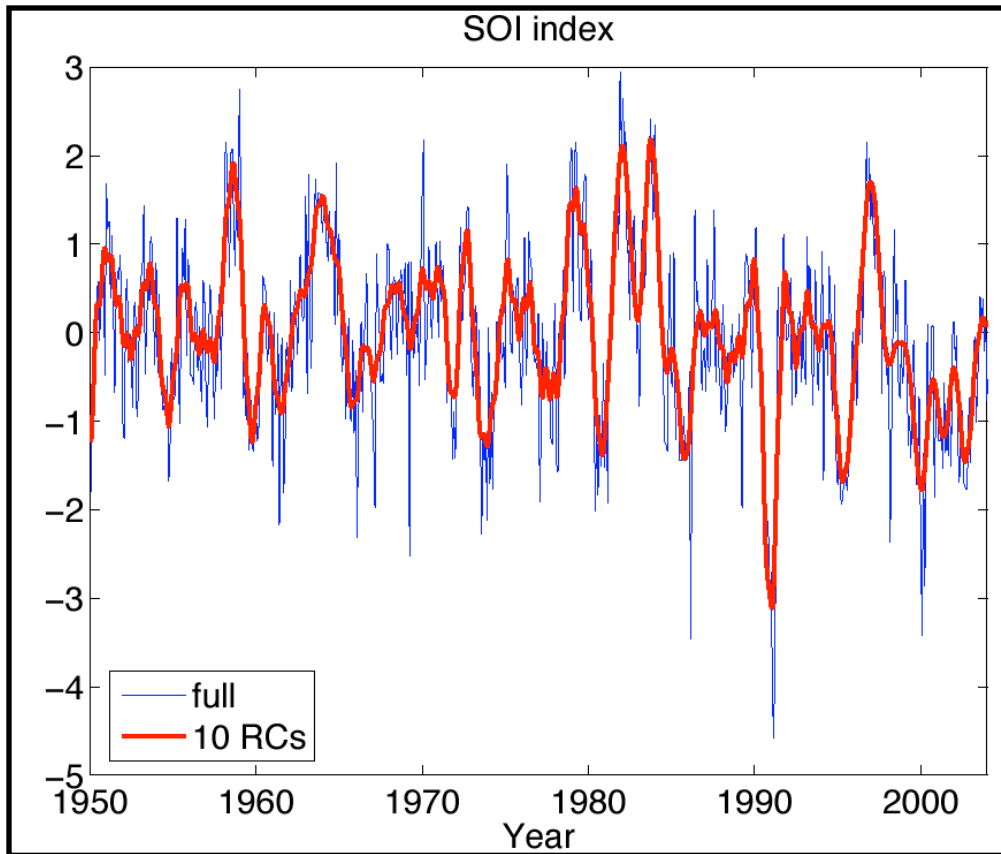


- SSA is good at isolating oscillatory behavior via paired eigenelements.
- SSA tends to lump signals that are longer-term than the window into
 - one or two trend components.

Selected References:

Vautard & Ghil (1989, *Physica D*);
Ghil *et al.* (2002, *Rev. Geophys.*) 12/28

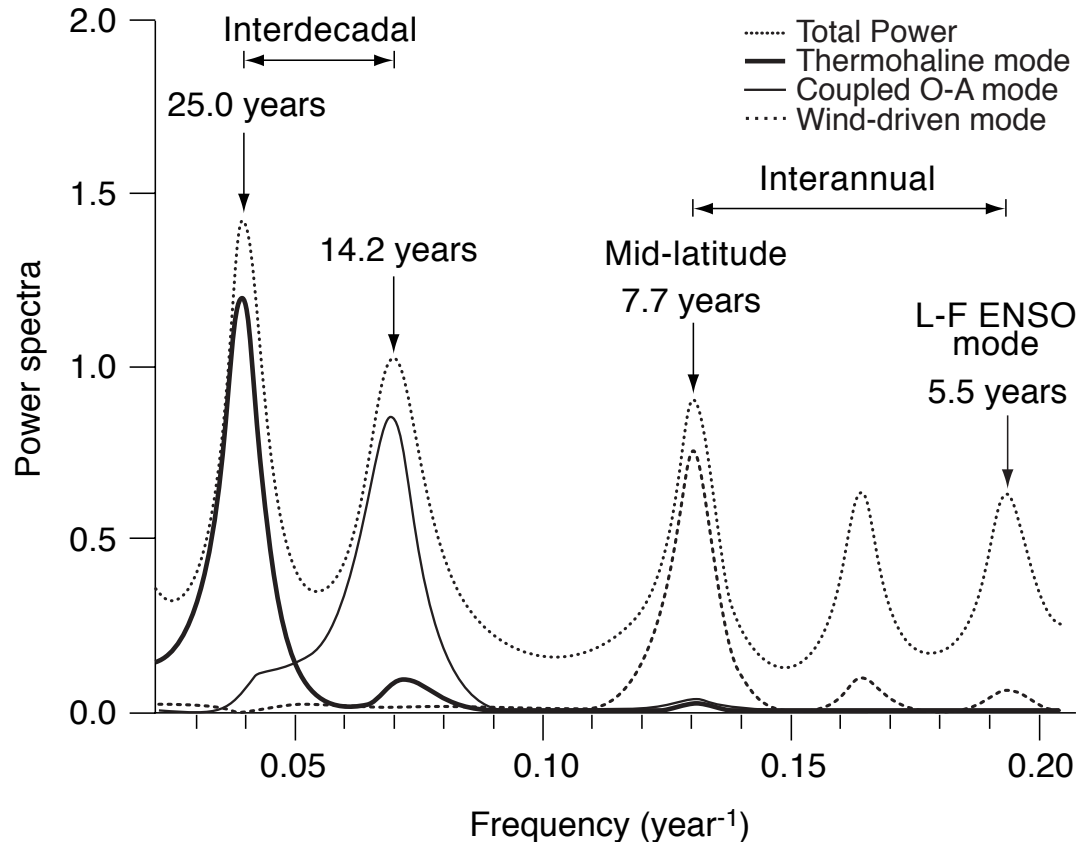
Singular Spectrum Analysis (SSA) and M-SSA (cont'd)



- Break in slope of SSA spectrum distinguishes “**significant**” from “**noise**” EOFs
- Formal Monte-Carlo test (Allen and Smith, 1994) identifies 4-yr and 2-yr ENSO oscillatory modes. A window size of $M = 60$ is enough to “resolve” these modes in a monthly SOI time series

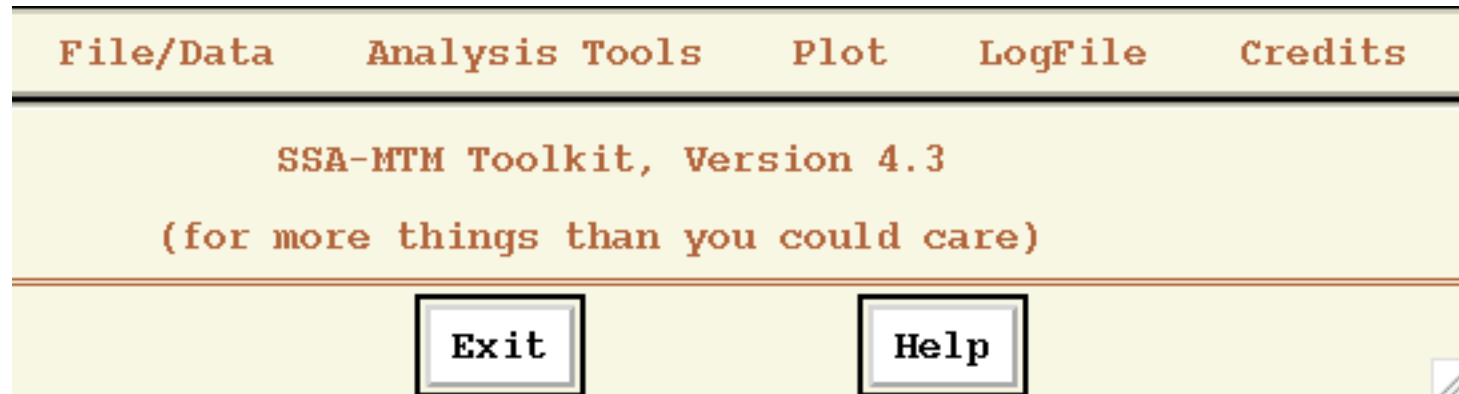
SSA (prefilter) + (low-order) MEM

o “Stack” spectrum

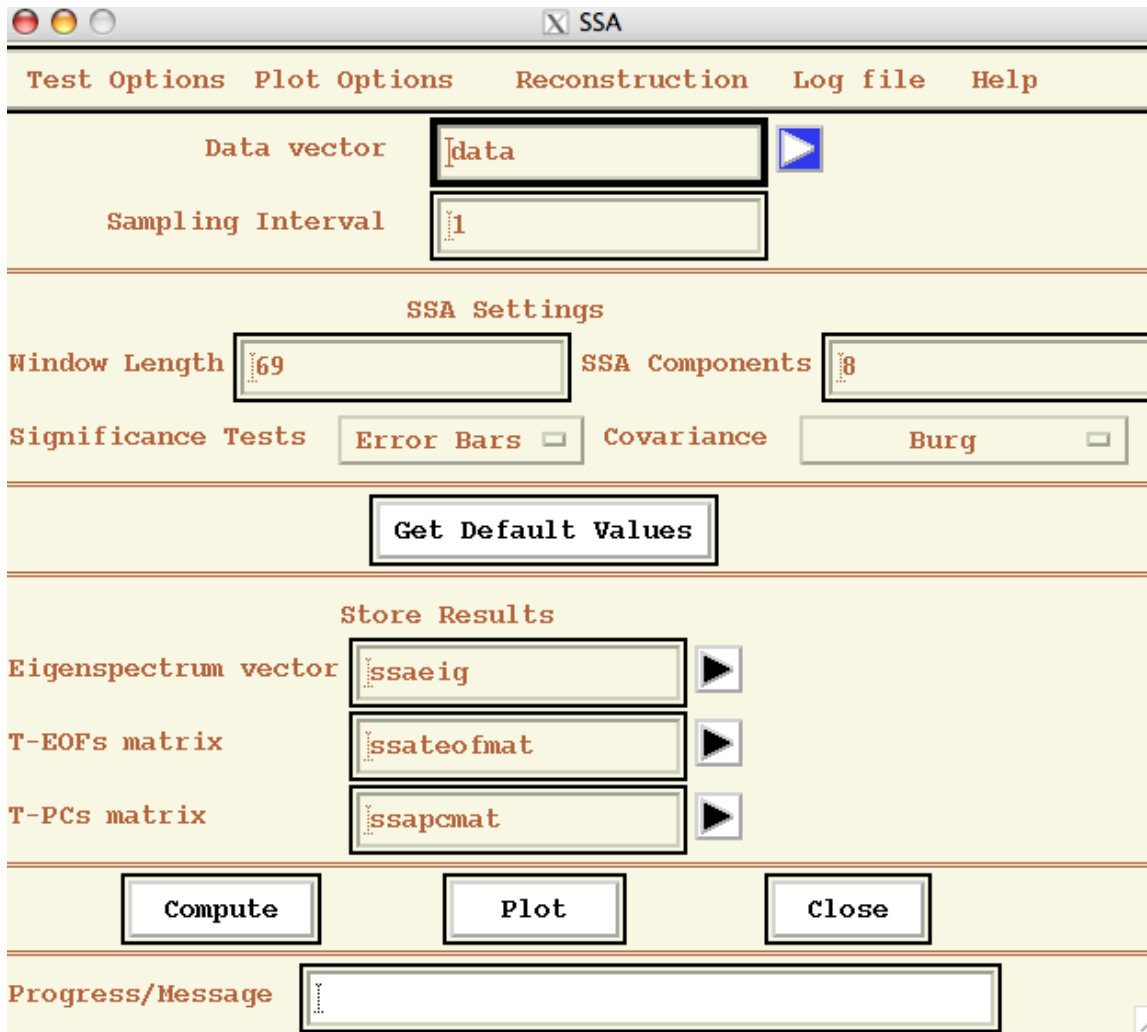


In good agreement with MTM peaks of **Ghil & Vautard (1991, *Nature*)** for the Jones *et al.* (1986) temperatures & stack spectra of Vautard *et al.* (1992, *Physica D*) for the IPCC “consensus” record (both global), to wit 26.3, 14.5, 9.6, 7.5 and 5.2 years.

Peaks at 27 & 14 years also in Koch sea-ice index off Iceland (Stocker & Mysak, 1992), etc.
Plaut, Ghil & Vautard (1995, *Science*)



- Ported to Sun, Dec, SGI, PC Linux, and Mac OS X
- Graphics support for [IDL](#) and [Grace](#)
- Precompiled binaries are available at www.atmos.ucla.edu/tcd/ssa
- Includes **Blackman-Tukey FFT**, **Maximum Entropy Method**, **Multi-Taper Method (MTM)**, **SSA and M-SSA**.
- Spectral estimation, decomposition, reconstruction & prediction.
- Significance tests of “**oscillatory modes**” vs. “**noise.**”



- **Free!!!**
- Data management with *named vectors & matrices*.
- *Default values* button.

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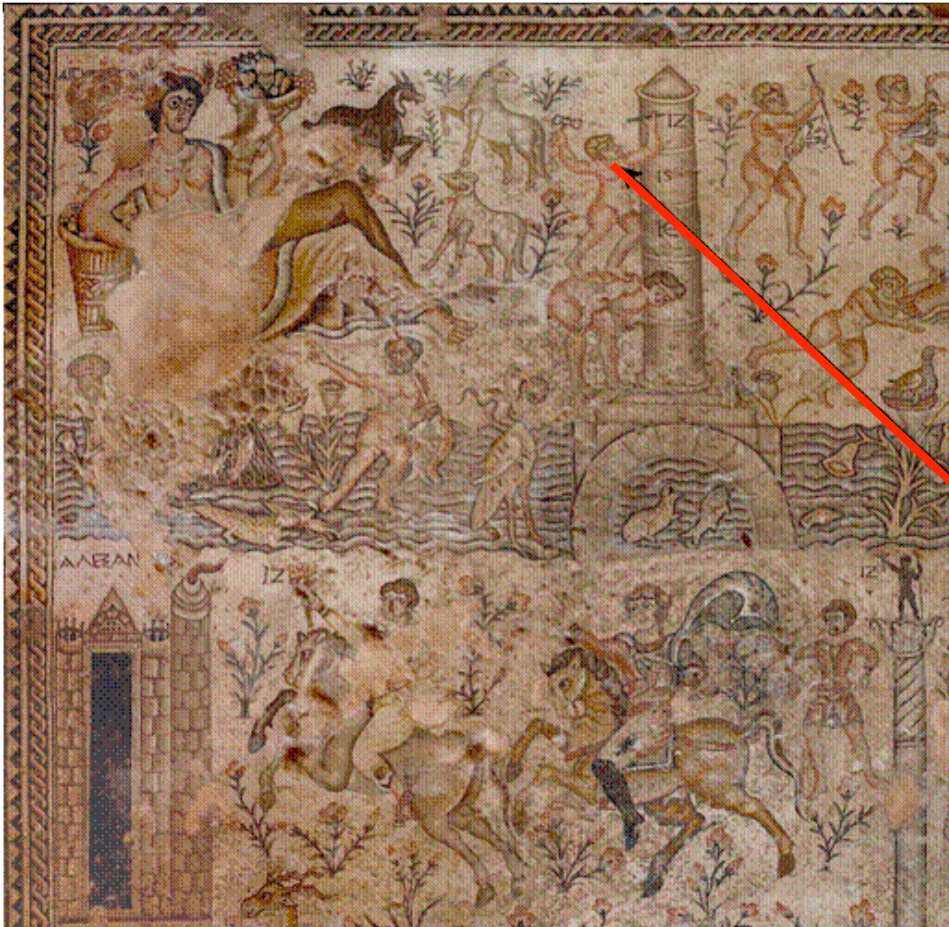
The Nile River Records Revisited: **How good were Joseph's predictions?**

Michael Ghil, ENS & UCLA

Yizhak Feliks, IIBR & UCLA,

Dmitri Kondrashov, UCLA

Why are there data missing?



Hard Work

- Byzantine-period mosaic from Zippori, the capital of Galilee (1st century B.C. to 4th century A.D.); photo by Yigal Feliks, with permission from the Israel Nature and Parks Protection Authority)

What to do about gaps?

- Most of the advanced *filling-in* methods are different flavors of **Optimal Interpolation (OI)**: Reynolds & Smith, 1994; Kaplan 1998).

Drawbacks: they either (i) require error statistics to be specified *a priori*; or (ii) derive it **only** from the interval of dense data coverage.

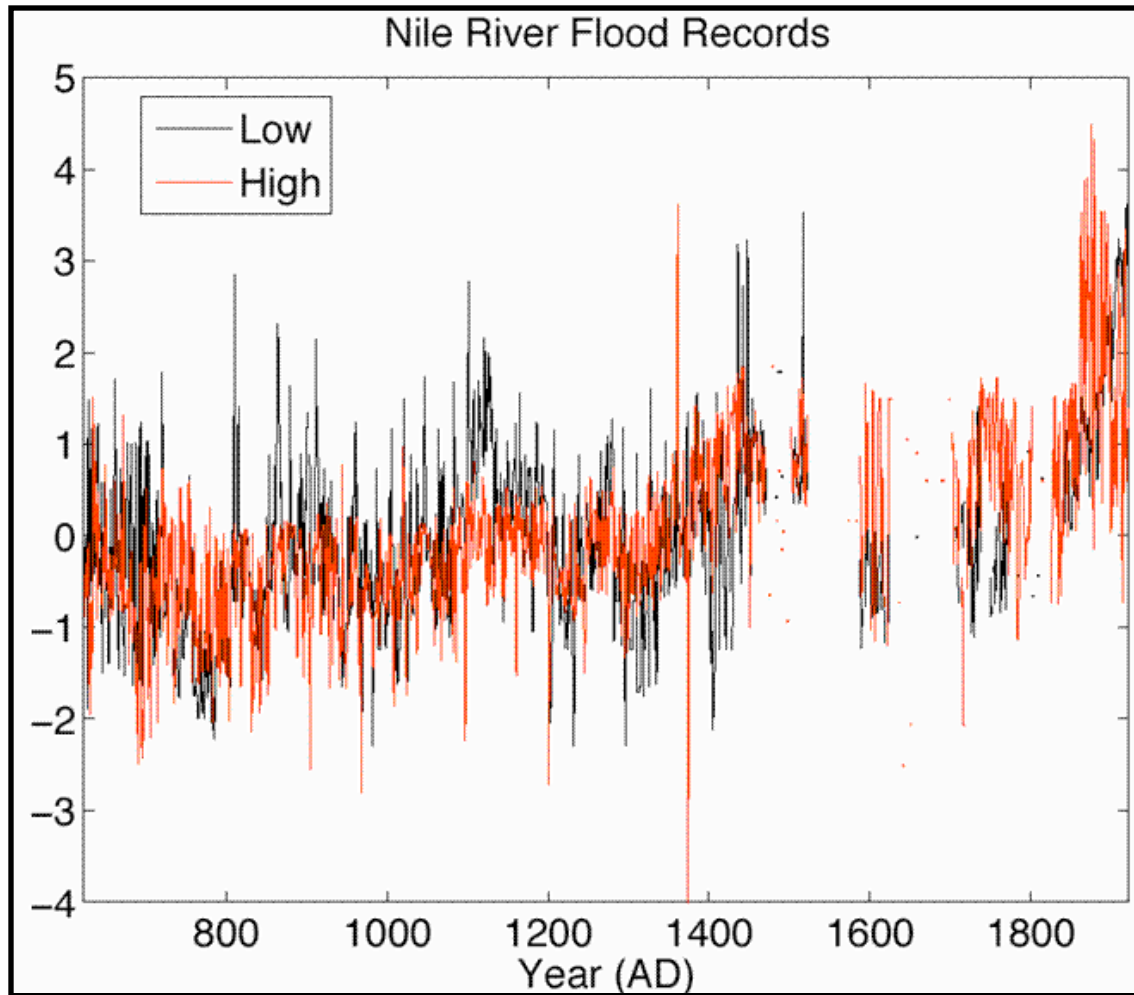
EOF Reconstruction (Beckers & Rixen, 2003): (i) iteratively compute **spatial-covariance** matrix using **all the data**; (ii) determine via cross-validation “**signal**” EOFs and use them to fill in the missing data; accuracy is similar to or better than **OI** (Alvera-Azcarate *et al.* 2004).

Drawbacks: uses **only** spatial correlations => cannot be applied to very **gappy** data.

We propose *filling in* gaps by applying **iterative SSA (or M-SSA):**

Utilize both spatial and temporal correlations of data => can be used for highly **gappy** data sets; simple and easy to implement!

Historical records are full of “gaps”....



Annual maxima and minima of the water level at the nilometer on Rodah Island, Cairo.

SSA (M-SSA) Gap Filling

Main idea: utilize **both spatial and temporal correlations** to iteratively compute self-consistent lag-covariance matrix; M-SSA with $M = 1$ is the same as the EOF reconstruction method of Beckers & Rixen (2003)

Goal: keep “**signal**” and truncate “**noise**” — usually a few leading EOFs correspond to the dominant oscillatory modes, while the rest is noise.

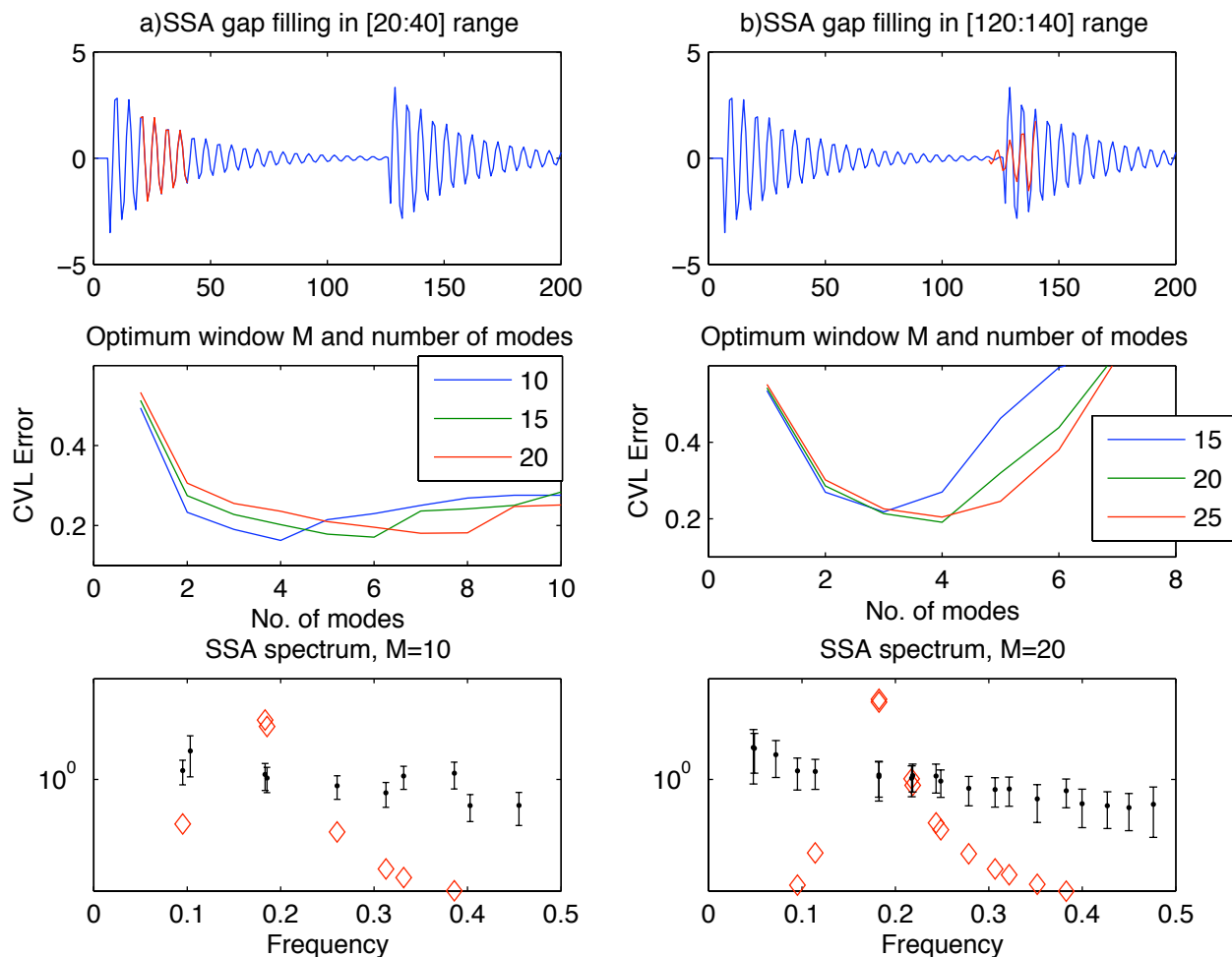
(1) for a given window width M : center the original data by computing the unbiased value of the mean and set the missing-data values to zero.

(2) start **iteration** with the **first EOF**, and replace the missing points with the reconstructed component (RC) of that EOF; **repeat the SSA algorithm** on the new time series, until convergence is achieved.

(3) repeat steps (1) and (2) with **two leading EOFs**, and so on.

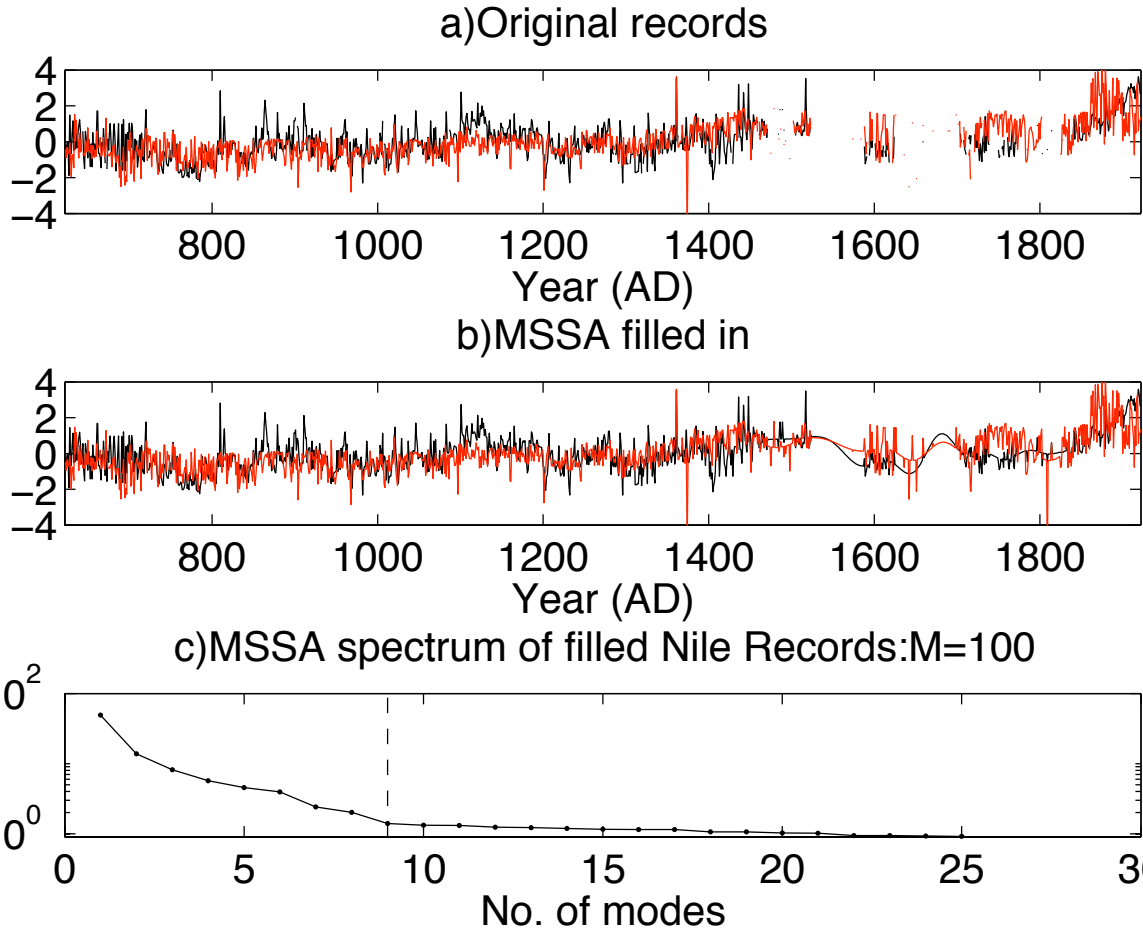
(4) apply **cross-validation** to optimize the value of M and the number of dominant SSA (M-SSA) modes K to fill the gaps: a portion of available data (selected at random) is flagged as missing and the RMS error in the reconstruction is computed.

Synthetic I: Gaps in Oscillatory Signal



- Very good gap filling for smooth modulation; OK for sudden modulation.

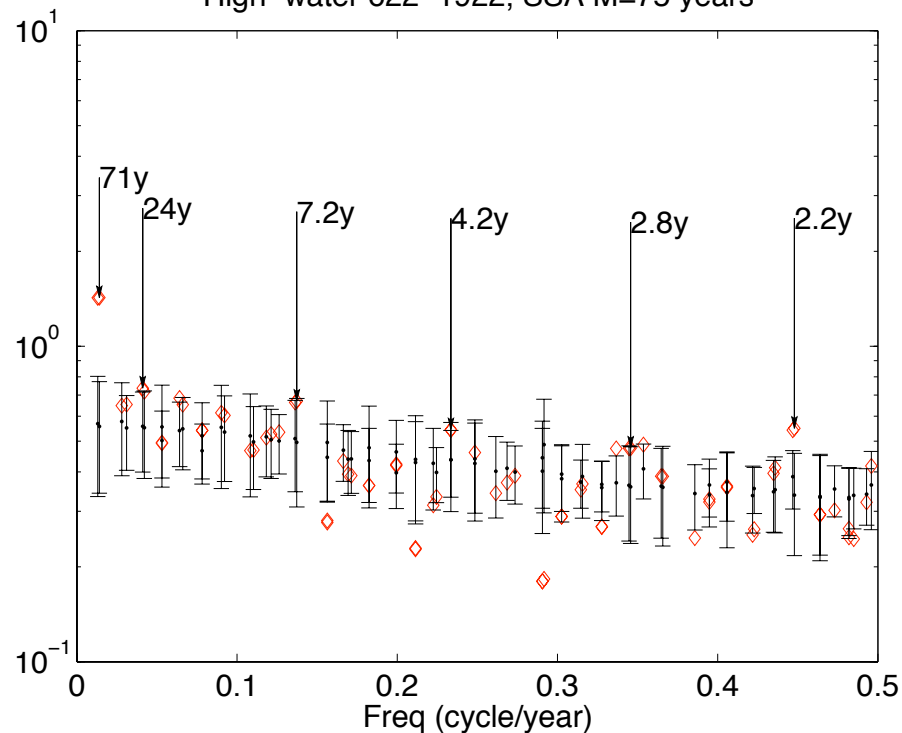
Nile River Records



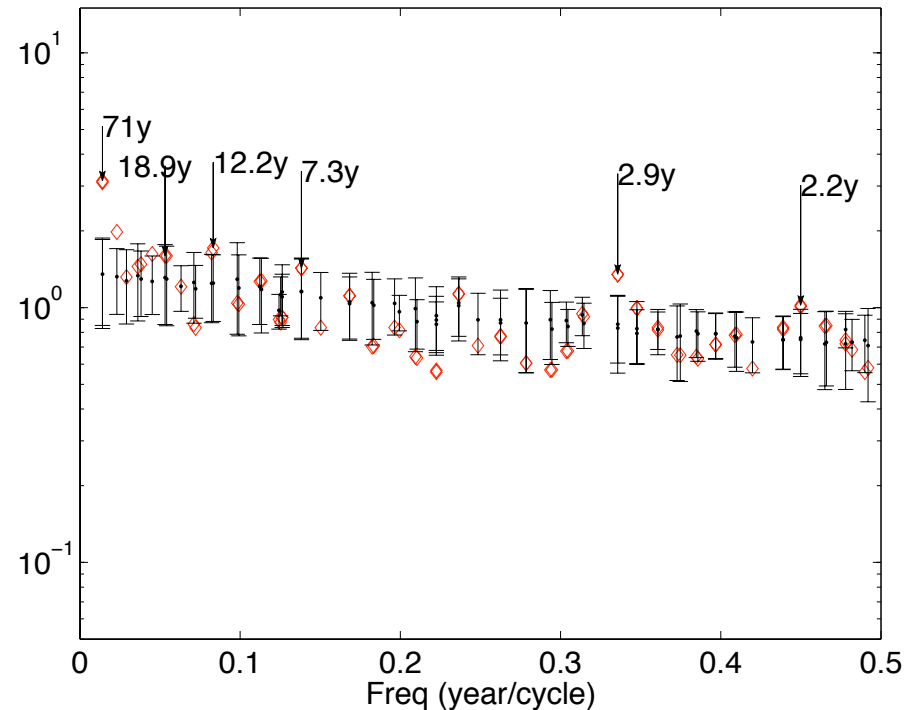
- High level —————
- Low level —————

MC-SSA of Filled-in Records

High-water 622–1922, SSA M=75 years



High-Low Water Difference, 622–1922, SSA M=75 years



SSA results for the extended Nile River records;
arrows mark highly significant peaks (at 95%), in both SSA and MTM.

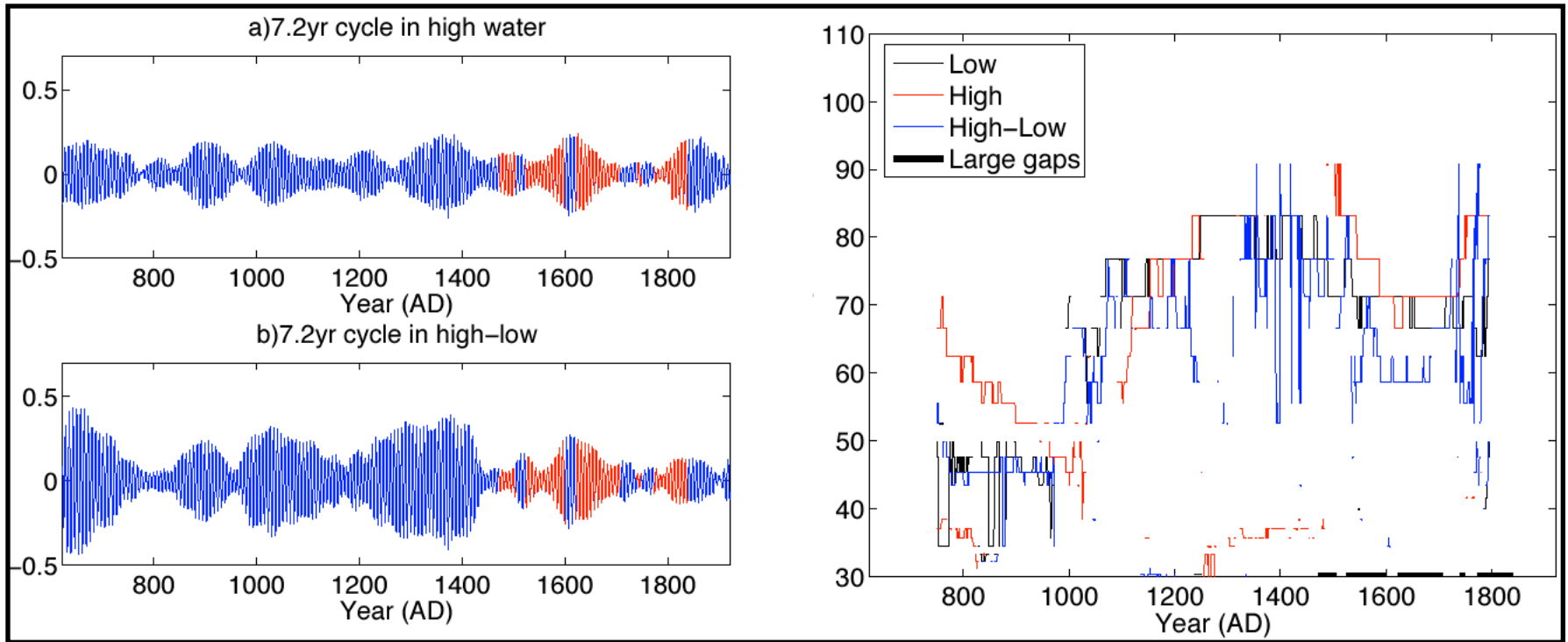
Table 1a: Significant oscillatory modes in short records (A.D. 622–1470)

Periods	Low	High	High-Low
40–100yr	64 (9.3%)	64 (6.9%)	64 (6.6%)
20–40yr		[32]	
10–20yr	12.2 (5.1%), 18.0 (6.7%)		12.2 (4.7%), 18.3 (5.0%)
5–10yr	6.2 (4.3%)	7.2 (4.4%)	7.3 (4.4%)
0–5yr	3.0 (2.9%), 2.2 (2.3%)	3.6 (3.6%), 2.9 (3.4%), 2.3 (3.1%)	2.9 (4.2%),

Table 1b: Significant oscillatory modes in extended records (A.D. 622–1922)

Periods	Low	High	High-Low
40–100yr	64 (13%)	85 (8.6%)	64 (8.2%)
20–40yr		23.2 (4.3%)	
10–20yr	[12], 19.7 (5.9%)		12.2 (4.3%), 18.3 (4.2%)
5–10yr	[6.2]	7.3 (4.0%)	7.3 (4.1%)
0–5yr	3.0 (4%), 2.2 (3.3%)	4.2 (3.3%), 2.9 (3.3%), 2.2 (2.9%)	[4.2], 2.9 (3.6%), 2.2 (2.6%)

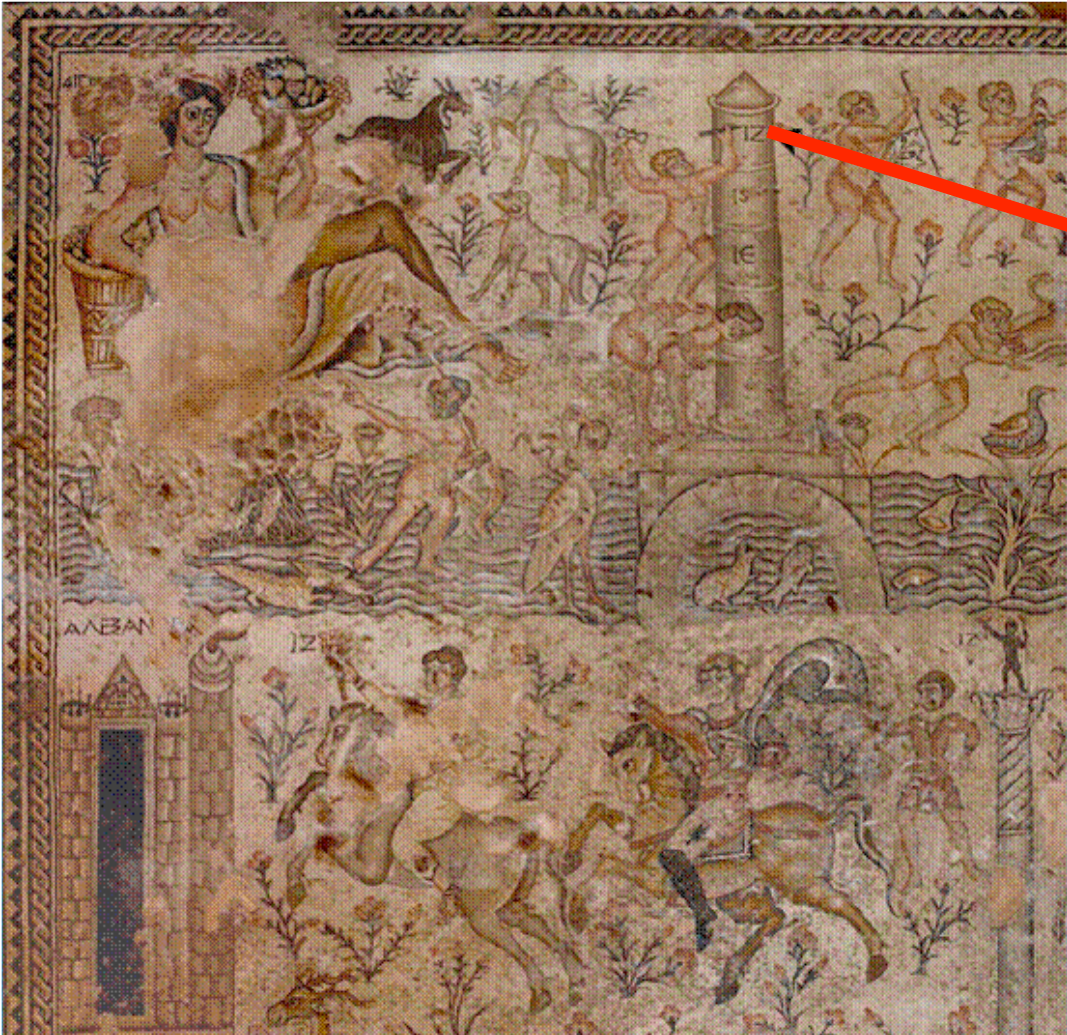
Significant Oscillatory Modes



SSA reconstruction of the 7.2-yr mode in the extended Nile River records:
(a) high-water, and (b) difference.
Normalized amplitude; reconstruction in the large gaps in red.

Instantaneous frequencies of the oscillatory pairs in the low-frequency range (40–100 yr).
The plots are based on multi-scale SSA [Yiou *et al.*, 2000]; local SSA performed in each window of width $W = 3M$, with $M = 85$ yr.

How good were Joseph's predictions?



Pretty good!

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Significance tests (“garde-fous”) in SSA

To check a spectral feature, e.g., **an oscillatory pair**:

1. Find pair for given **data set** $\{X_n: n = 1, 2, \dots, N\}$ and **window width** M .
2. Apply **statistical significance tests** (MC-SSA, etc.).
3. Check **robustness** of pair by changing M , sampling interval τ_s , etc.
3. Apply **additional methods** (MTM, wavelets, etc.) and their tests to $\{X_n\}$.
4. Obtain **additional time series** pertinent to the same phenomenon $\{Y_m\}$, etc.
5. Apply steps (1)–(3) to these data sets.
6. Use **multi-channel SSA (M-SSA)** and other multivariate methods to check mutual dependence between $\{X_n\}$, $\{Y_m\}$, etc.
7. Based on steps (1)–(6), try to provide a **physical explanation** of the mode.
8. Use (7) to **predict an as-yet-unobserved feature** of the data sets.
9. If this new feature **is found in new data**, go on to next problem.
10. **If not, go back** to an earlier step of this list.

(*) **Ghil, M.**, M. R. Allen, M. D. Dettinger, K. Ide, D. Kondrashov, M. E. Mann, A. W. Robertson, A. Saunders, Y. Tian, F. Varadi, and P. Yiou, 2002: Advanced spectral methods for climatic time series, *Rev. Geophys.*, **40**(1), pp. 3.1–3.41, doi: 10.1029/2000RG000092.

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Reserve slides

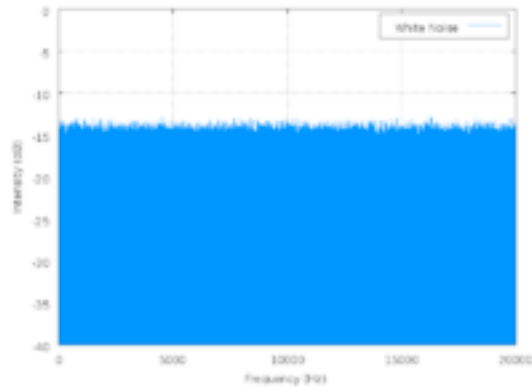
Diapos de réserve

Singular

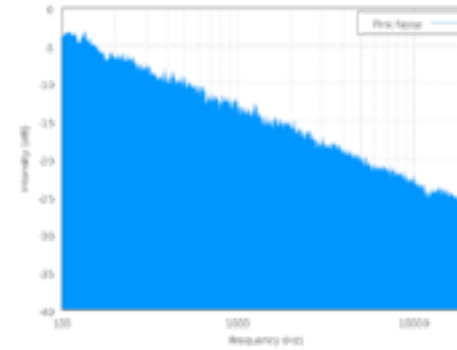
Spectrum

Analysis

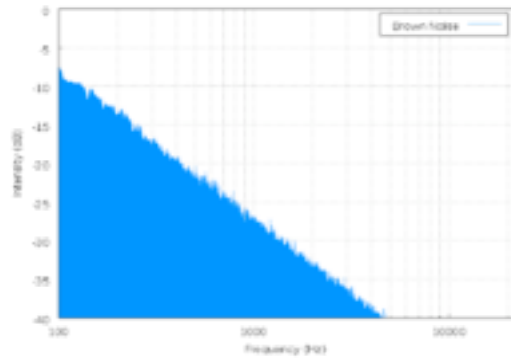
Noise “colors”



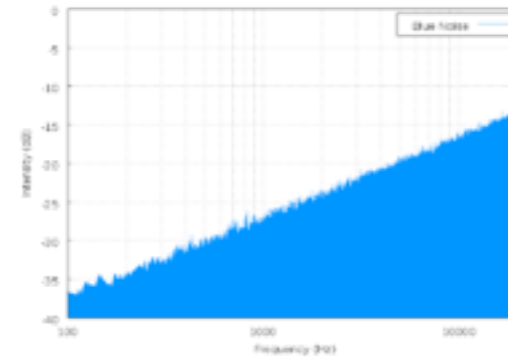
White noise, $S \sim f^0$



Pink (or $1/f$) noise, $S \sim f^{-1}$



Red (or Brown) noise, $S \sim f^{-2}$



Blue noise, $S \sim f^{+1}$

SSA for Southern Oscillation Index (SOI)

SOI = mean monthly values of Δp_s (Tahiti – Darwin)

Results (“undigested”) from 1933–1988 time interval (*)

1. For $18 < M < 60$ months, singular spectra show a clear break at $5 < S < 17$ (= “deterministic” part; $M - S =$ “noise”);
2. 3 pairs of EOFs stand out:
EOFs 1 + 2 (27%), 3 + 4 (19.7%), and 9 + 10 (3%);
3. the associated periods are ~
60 mos. (“ENSO”), 30 mos. (QBO”), and 5.5 mos. (?!)

(*) E. M. (“Gene”) Rasmusson, X. Wang, and C.F. Ropelewski, 1990:
The biennial component of ENSO variability. *J. Marine Syst.*, **1**, 71–96.

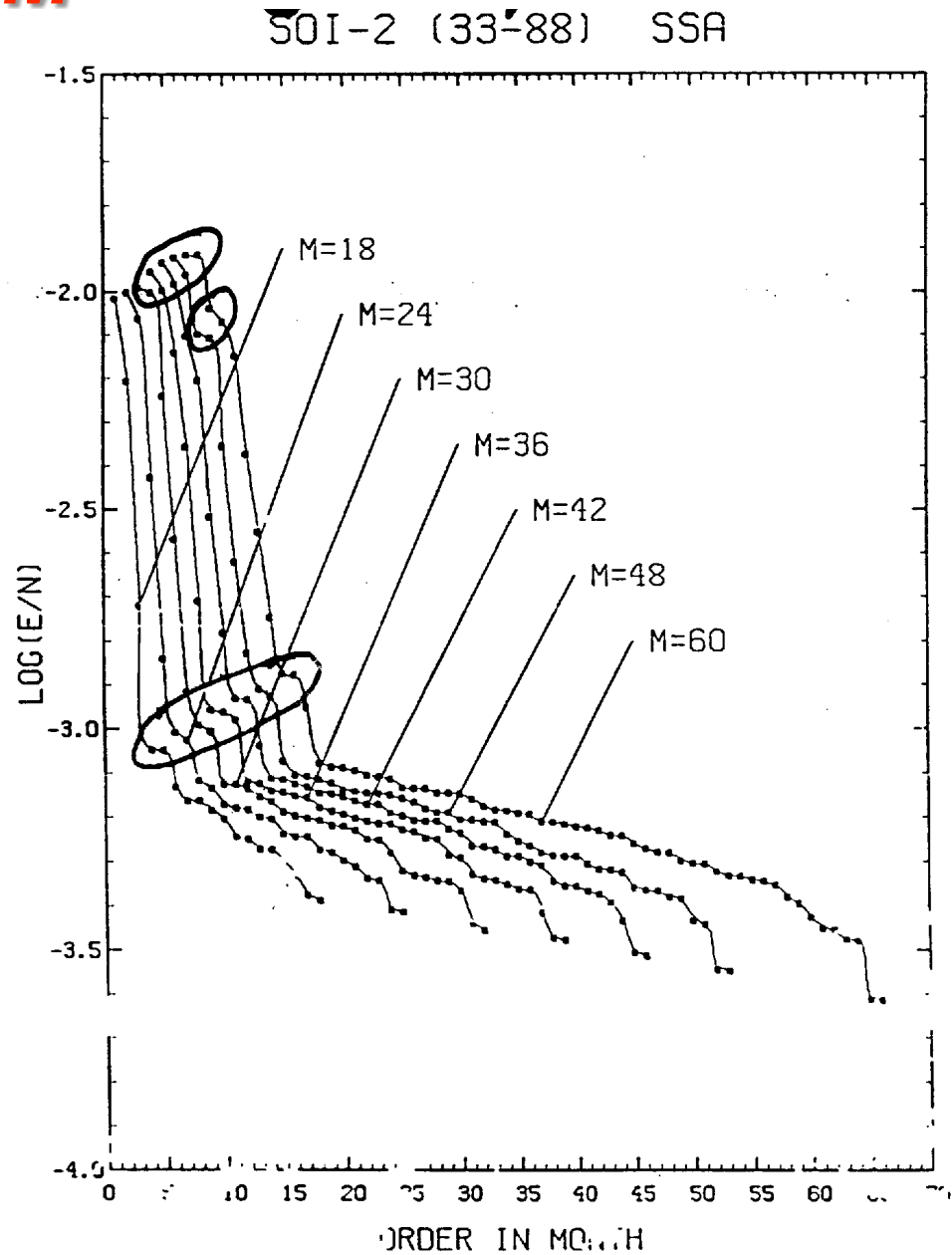
Variable window size M

Sampling interval – $\tau_s = 1$ month

Window width $M\tau_s$:

$18\tau_s < \tau_w < 60\tau_s$ or

$1.5 \text{ yr} < \tau_w < 5 \text{ yr}$.



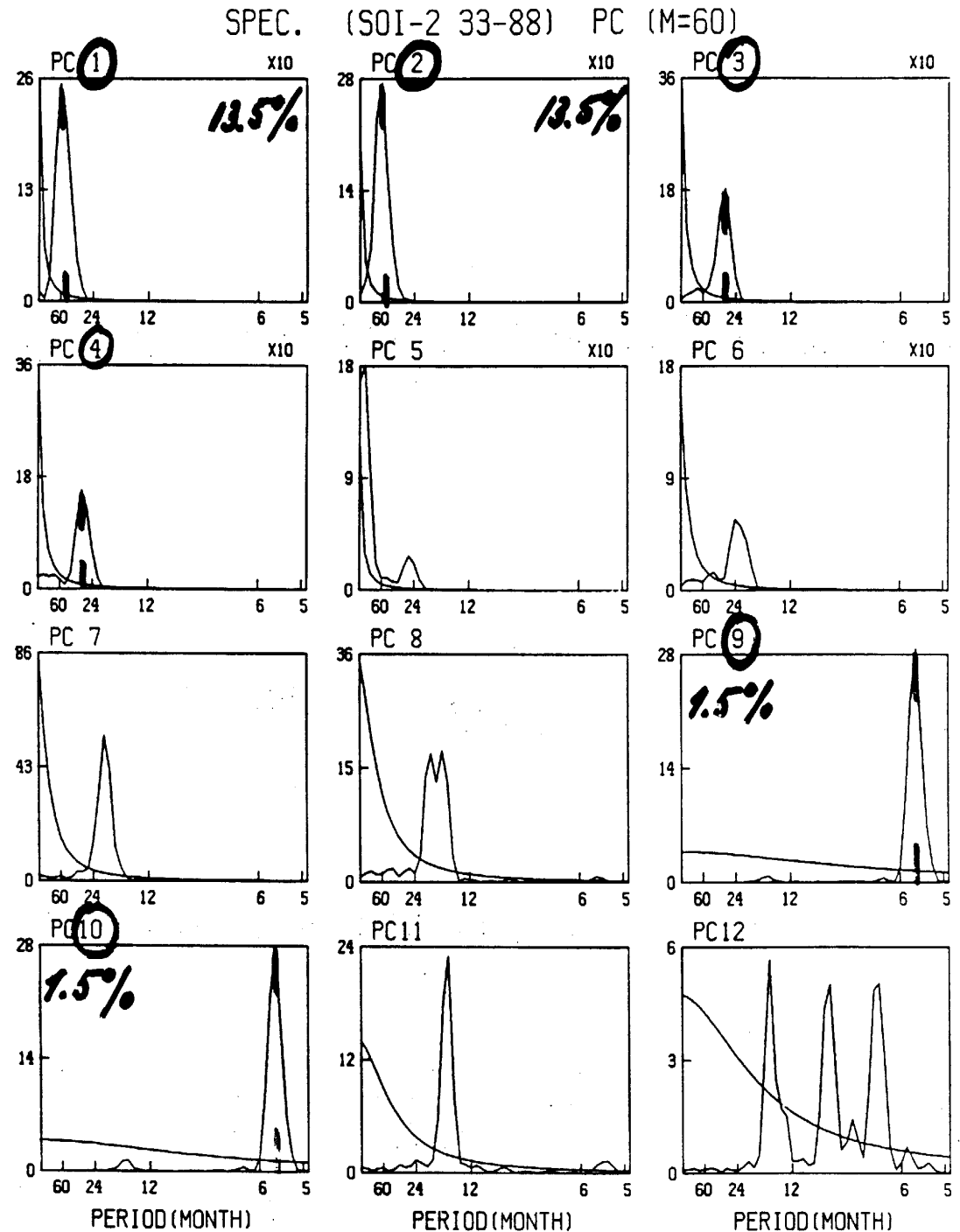
Spectral peaks ($M = 60$)

Each principal component (PC) is Fourier analyzed separately; individual variance indicated as well.

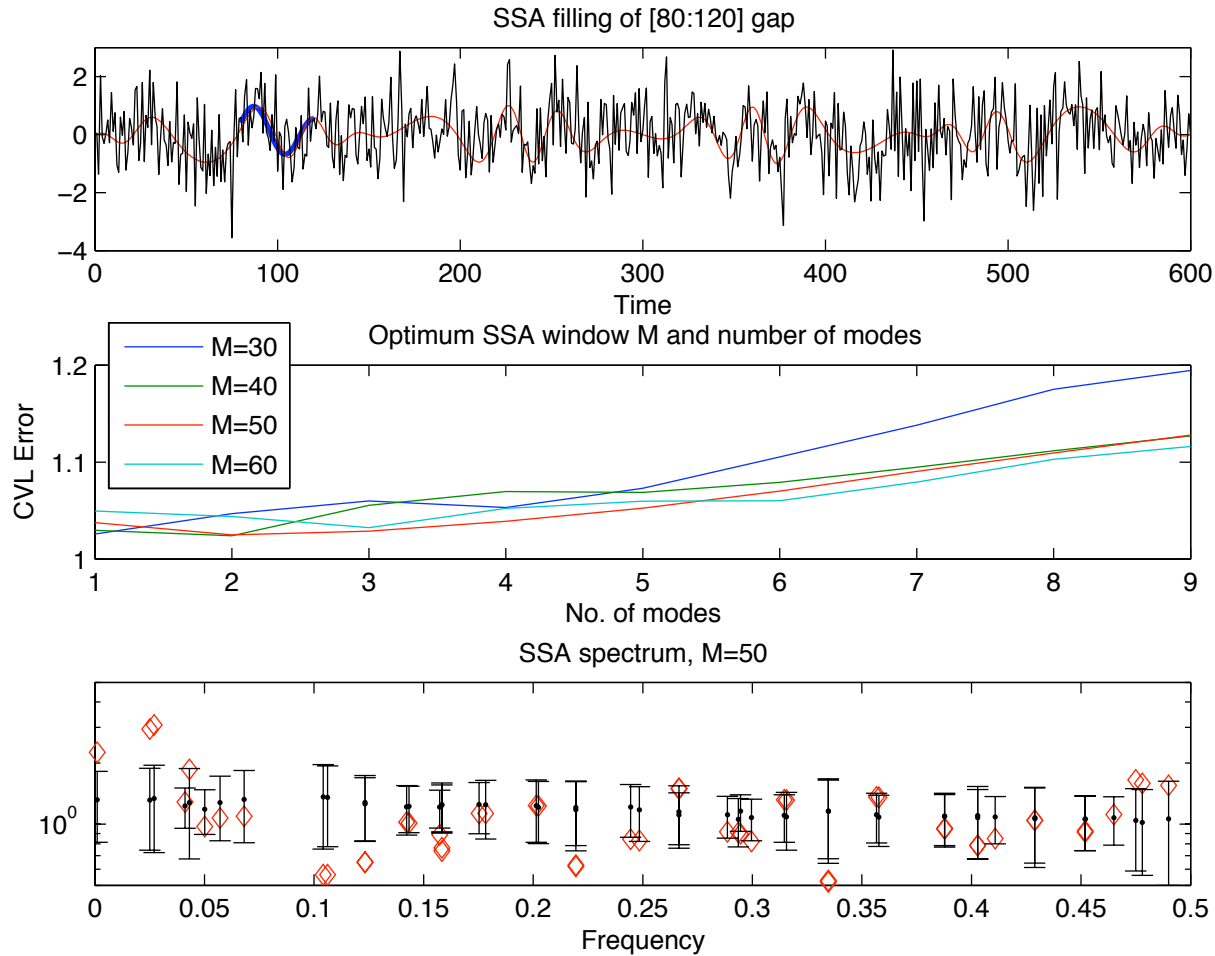
PCs (1+2) – period = 60 months, low-frequency or “ENSO” or quasi-quadrennial (QQ) component;

PCs (3+4) – period = 30 months quasi-biennial (QB) component;

PCs (9+10) – period = 5.5 months

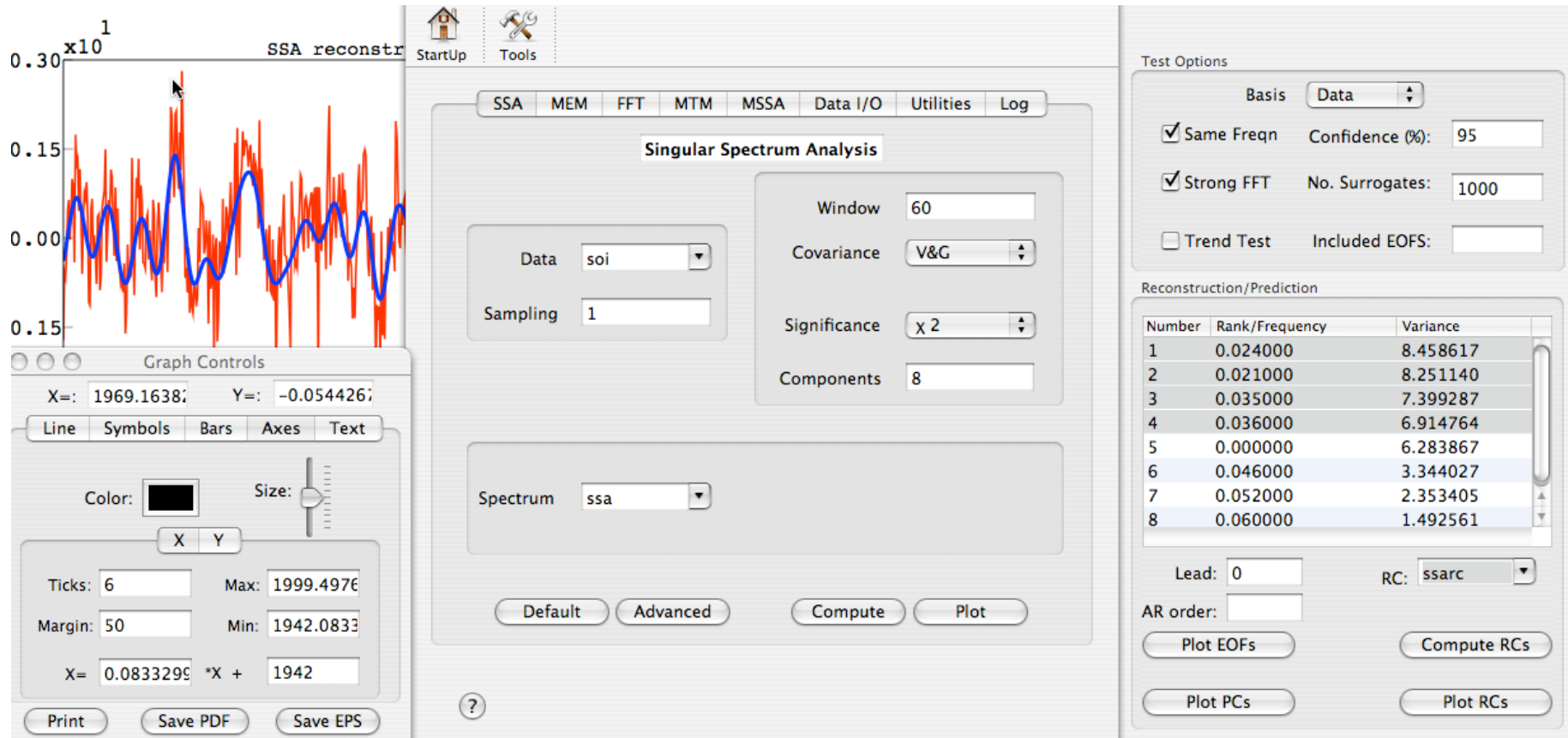


Synthetic II: Gaps in Oscillatory Signal + Noise



$$x(t) = \sin\left(\frac{2\pi}{300}t\right) * \cos\left(\frac{2\pi}{40}t + \frac{\pi}{2}\sin\frac{2\pi}{120}t\right)$$

kSpectra Toolkit for Mac OS X



• \$\$... but: *Project files, Automator WorkFlows, Spotlight* and more!

• www.spectraworks.com

Un peu de bibliographie

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- <http://www.atmos.ucla.edu/tcd/ssa> .
- <http://www.r-project.org> .

The Nile River Basin initiative will greatly modify the flow along the longest & best-documented river system in the world ...

